

1 **Multi-year Predictions of North Atlantic Hurricane Frequency:**
2 **Promise and limitations**

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14 Submitted to *J. Climate*
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19 **MANUSCRIPT IN PRESS.**
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22 Submitted: July 19, 2012

23 Revised: December 1, 2012

24 Accepted: February 1, 2013
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Abstract

Retrospective predictions of multi-year North Atlantic hurricane frequency are explored, by applying a hybrid statistical-dynamical forecast system to initialized and non-initialized multi-year forecasts of tropical Atlantic and tropical mean sea surface temperatures (SSTs) from two global climate model forecast systems. By accounting for impacts of initialization and radiative forcing, retrospective predictions of five-year mean and nine-year mean tropical Atlantic hurricane frequency show significant correlation relative to a null hypothesis of zero correlation. The retrospective correlations are increased in a two-model average forecast and by using a lagged-ensemble approach, with the two-model ensemble decadal forecasts hurricane frequency over 1961-2011 yielding correlation coefficients that approach 0.9.

These encouraging retrospective multi-year hurricane predictions, however, should be interpreted with care: although initialized forecasts have higher nominal skill than uninitialized ones, the relatively short record and large autocorrelation of the time series limits our confidence in distinguishing between the skill due to external forcing and that added by initialization. The nominal increase in correlation in the initialized forecasts relative to the uninitialized experiments is due to improved representation of the multi-year tropical Atlantic SST anomalies. The skill in the initialized forecasts comes in large part from the persistence of a mid-1990s shift by the initialized forecasts, rather than from predicting its evolution. Predicting shifts like that observed in 1994-1995 remains a critical issue for the success of multi-year forecasts of Atlantic hurricane frequency. The retrospective forecasts highlight the possibility that changes in observing system impact forecast performance.

1

2 **I. Introduction**

3 Predicting and projecting future North Atlantic hurricane activity is a topic of
4 scientific interest (*e.g.*, Gray 1984; Knutson and Tuleya 2004; Emanuel 2005; Camargo
5 *et al.* 2007a; Vecchi *et al.* 2008; Smith *et al.* 2010; Knutson *et al.* 2010; Vecchi *et al.*
6 2011; Villarini *et al.* 2011.a; Villarini and Vecchi 2012b-d) and high societal significance
7 (Pielke Jr. *et al.* 2008; Mendelsohn *et al.* 2012; Peduzzi *et al.* 2012). Seasonal basin-wide
8 frequency of North Atlantic hurricanes has exhibited variability on a variety of
9 timescales, from interannual to multi-decadal, although it remains unclear whether there
10 has been any century-scale trend in Atlantic hurricane frequency (*e.g.*, Mann and
11 Emanuel 2006; Vecchi and Knutson 2008, 2011; Landsea *et al.* 2011; Villarini *et al.*
12 2011b).

13 The scientific basis for predictions of seasonal hurricane activity at leads of one to
14 three seasons has been developed (*e.g.*, Gray 1984; Elsner and Jagger 2006; Vitart 2006;
15 Camargo *et al.* 2007a,b; Vitart *et al.* 2007; Klotzbach and Gray 2009; Wang *et al.* 2009;
16 Kim and Webster 2010; LaRow *et al.* 2010; Zhao *et al.* 2010; Alessandri *et al.* 2011;
17 Chen and Lin 2011; Vecchi *et al.* 2011; Villarini and Vecchi 2012d), leading to the
18 identification of different potential sources of skill, both local and remote.

19 Decadal to centennial projections of seasonal hurricane activity in response to
20 changes in external forcing (greenhouse gases, aerosols, volcanoes, and solar) have been
21 made (*e.g.* Oouchi *et al.* 2006; Knutson *et al.* 2008; Emanuel *et al.* 2008; Gualdi *et al.*
22 2008; Vecchi *et al.* 2008; Sugi *et al.* 2009, 2012; Zhao *et al.* 2009; Bender *et al.* 2010;
23 Knutson *et al.* 2010; Knutson *et al.* 2010; Villarini *et al.* 2011a; Zhao and Held 2011;

Villarini and Vecchi 2012b,c). The basis for these projections is the possibility that radiatively-forced climate change could influence the climatic conditions to which hurricanes are sensitive, such as large-scale circulation, wind shear, ocean temperatures, potential intensity and humidity (*e.g.*, Emanuel 1987, 2007; Broccoli and Manabe 1990; Shen *et al.* 2000; Knutson and Tuleya 2004; Camargo *et al.* 2007b; Vecchi and Soden 2007a,b). Recent model results span a relatively wide range of possibilities for North Atlantic hurricane frequency (including increases or decreases) under enhanced CO₂-induced warming, while there is a wider tendency for hurricane intensity to increase in these studies (*e.g.*, Knutson and Tuleya 2004; Knutson *et al.* 2008, Emanuel *et al.* 2008, Gualdi *et al.* 2008; Knutson *et al.* 2008; Vecchi *et al.* 2008; Sugi *et al.* 2009, 2012; Zhao *et al.* 2009, Bender *et al.* 2010; Knutson *et al.* 2010, Villarini *et al.* 2011a; Villarini and Vecchi 2012b,c). There are indications that changes in atmospheric aerosols could influence past and projected hurricane activity, with increases (decreases) in Atlantic aerosol loading driving decreases (increases) in Atlantic hurricane activity (Mann and Emanuel 2006; Evan *et al.* 2009, Villarini and Vecchi 2012b,c).

Assessing hurricane predictability at intermediate timescales, between seasonal predictions and multi-decadal projections, is an emerging field of research. In addition to potential influences from changes in radiative forcing, internal variations of the climate system could play a large role in changes of hurricane frequency on timescales of decades (*e.g.*, Goldenberg *et al.* 1996; Zhang and Delworth 2006, 2009; Knight *et al.* 2006; Latif *et al.* 2007; Dunstone *et al.* 2011; Villarini *et al.* 2011; Villarini and Vecchi 2012b). There are physical reasons to expect coherent multi-year hurricane variations to be tied to ocean changes (*e.g.*, Goldenberg *et al.* 1996, Zhang and Delworth 2005, 2006, 2009;

1 Knight *et al.* 2006, Latif *et al.* 2007, Dunstone *et al.* 2011). There is also indication that
2 some of the relevant ocean changes may be potentially predictable on decadal timescales
3 (*e.g.*, Griffies and Bryan 1997a,b; Pohlmann *et al.* 2004; Collins *et al.* 2006; Pohlmann *et*
4 *al.* 2009; Msadek *et al.* 2010; Smith *et al.* 2010; Teng *et al.* 2011; Chikamoto *et al.* 2012;
5 van Oldenborgh *et al.* 2012; Rosati *et al.* 2012; Yang *et al.* 2012; Yeager *et al.* 2012). As
6 decadal variability and the associated predictability can result from both internal and
7 externally forced fluctuations (*e.g.*, Rotstayn and Lohmann 2002; Hawkins and Sutton
8 2009; Chang *et al.* 2011a; Villarini *et al.* 2011; Booth *et al.* 2012; Villarini and Vecchi
9 2012b), one has to consider skill arising from both external factors and internal variability
10 on multi-year timescales. A number of modeling groups are now following the same
11 framework for the Fifth Coupled Model Intercomparison Project (CMIP5; Taylor *et al.*
12 2012) to be assessed as part of the 5th Assessment Report of the Intergovernmental Panel
13 on Climate Change (IPCC-AR5), by performing decadal predictions initialized with
14 estimates of the observed state of the climate system (Taylor *et al.* 2012, Meehl *et al.*
15 2012). While for sea surface temperature (SST), most of the skill on multi-year
16 timescales arises from predicting the warming trend associated with radiative forcing
17 changes (*e.g.*, van Oldenborgh *et al.* 2012; Rosati *et al.* 2012), there is at least one study
18 suggesting that initialization can increase the skill in multi-year hurricane forecasts
19 (Smith *et al.* 2010; henceforth S10). In this paper we explore the ability of a hybrid
20 statistical-dynamical hurricane forecasting system to retrospectively predict multi-year
21 hurricane activity in the Atlantic using two different coupled climate models, including
22 the one used by S10. We explore the skill of North Atlantic hurricane frequency resulting
23 from changing radiative forcing and from natural variability. We assess the improvement

1 in skill due to initialization and discuss the source of this improved skill and its
2 implications for future multi-year forecasts of North Atlantic hurricane frequency.

3 **II. Data and Methods**

4 *A. Statistical hurricane emulator:*

5 We use a hybrid statistical-dynamical North Atlantic hurricane frequency prediction
6 framework to explore the predictability of multi-year hurricane activity. This framework
7 has been shown to exhibit retrospective skill in seasonal hurricane forecasts from as early
8 as boreal winter prior to the hurricane season (Vecchi *et al.* 2011). It combines a
9 statistical emulator of a high-resolution dynamical atmospheric model (Zhao *et al.* 2009,
10 2010) and initialized forecasts of SST. The statistical emulator is formulated as a Poisson
11 regression model with two predictors: Tropical Atlantic SST and Tropical-mean SST,
12 each averaged over the August-October season.

13 The choice of these two predictors is motivated by dynamical considerations,
14 observed relationships between hurricane activity and SST, and the sensitivity of
15 dynamical models to SST perturbations. Observational analyses have highlighted
16 correlations between SST changes in the tropical Atlantic and hurricane activity indices
17 (*e.g.*, Elsner and Jagger 2006; Emanuel 2005). However, observational correlations as
18 high or higher have been found between hurricane activity and the weighted difference
19 between Atlantic and tropical-mean SSTs (the SST changes in the Atlantic relative to the
20 tropics, or “Relative SST”) by other studies (*e.g.*, Swanson 2007, 2008; Vecchi *et al.*
21 2008; Villarini *et al.* 2010, 2011.a, 2012; Villarini and Vecchi 2012). The physical basis
22 for exploring relative SST as a predictor of hurricane activity is based on the tendency of
23 free tropospheric temperature changes to follow those of tropical-mean SST (Sobel *et al.*

2002) or SSTs in the Indo-Pacific region where the bulk of tropical convection resides (Tang and Neelin 2004) as described by the Weak Temperature Gradient approximation (Sobel and Bretherton 2000). An Atlantic SST warming that is larger than that of the tropical average, with a tropospheric warming in the Atlantic that follows tropical-mean SST, would lead to a large-scale destabilization of the atmosphere in the Atlantic, to changes in the large-scale vorticity, shear and atmospheric humidity, as well as to increases in TC potential intensity (*e.g.*, Latif *et al.* 2007; Vecchi and Soden 2007; Gualdi *et al.* 2008; Sugi *et al.* 2009, 2012; Zhao *et al.* 2009; Xie *et al.* 2010; Zhao and Held 2011; Ramsay and Sobel 2011; Camargo *et al.* 2012; Vecchi *et al.* 2012). Supporting the notion of relative SST as a predictor for Atlantic hurricane activity, dynamical modeling studies have found that the threshold for TC genesis under projected climate changes over the 21st century increases along with the overall tropical warming (*e.g.* Knutson *et al.* 2008). The interannual, decadal and climate change response of North Atlantic TC frequency simulated with a across a range of dynamical frameworks is also well explained by relative SST (*e.g.*, Vecchi *et al.* 2008; Sugi *et al.* 2009, 2012; Zhao *et al.* 2009, 2010; Vecchi *et al.* 2011; Villarini *et al.* 2011.a; Knuston *et al.* 2012; Zhao and Held 2012), although strong departures from moist adiabatic warming can complicate relative SST models of hurricane frequency (*e.g.*, Vecchi *et al.* 2012).

Following Vecchi *et al.* (2011), we model the rate of occurrence (λ ; the expected value of the aggregate seasonal number) of North Atlantic hurricane frequency using a Poisson regression model as follows:

$$\lambda = e^{1.707 + 1.388SST_{MDR} - 1.521SST_{TROP}} \quad (\text{Eq.1})$$

1 where SST_{MDR} and SST_{TROP} are anomalies in the regional SST indices relative to the
 2 1982–2005 average, as described in Section II.C. SST_{MDR} is the average over the
 3 hurricane main development region (80°W–20°W, 10°N–25°N), and SST_{TROP} is the
 4 global, 30°S–30°N average of SST. As discussed in Vecchi *et al.* (2011), this statistical
 5 emulator of the sensitivity of hurricane frequency to SST changes in the Zhao *et al.*
 6 (2009, 2010) high-resolution atmospheric model was trained across a broad range of
 7 climate states, including multiple realizations of the historical period and various
 8 projections of 21st century SST change. This statistical model was trained against a wide
 9 range of climate states, and its performance against the observed record satisfies a
 10 necessary condition for its application to interannual to decadal prediction (Vecchi *et al.*
 11 2011). The parameters in this statistical emulator, built on the output of a high-resolution
 12 AGCM, are very similar to those that arise from modeling adjusted hurricane frequency
 13 over the 1878–2008 period (Villarini *et al.* 2012). This statistical emulator is able to
 14 reproduce much of the observed variability in hurricane activity ($r^2=0.58$; Vecchi *et al.*
 15 2011), and its ability to recover changes in hurricane frequency compares well with
 16 hindcasts and projections from high-resolution dynamical models (*e.g.*, Zhao *et al.* 2009,
 17 2010; Villarini *et al.* 2011a; Knutson *et al.* 2012). The low computational cost of the
 18 statistical emulator allows us to efficiently perform a variety of retrospective forecasts
 19 using multiple input datasets, described below.

20 *B. Global climate model predictions:*

21 The statistical emulator (described above) is applied to predictions of SST from two
 22 global climate models: NOAA Geophysical Fluid Dynamics Laboratory (GFDL) CM2.1
 23 and UKMetOffice (UKMO) Decadal Prediction System (DePreSys) Perturbed Physics

Ensemble (PPE), referred to as GFDL-DecPre and UKMO-DePreSys, respectively. The forecast system specifications are summarized in Table 1. These models are just two of those what will be part of the CMIP5 decadal prediction experiments, although the CMIP5 version of UKMO-DePreSys is slightly different from the one used here. Exploration of those models allows us to compare the behavior of a prediction system that has shown skill in interannual hurricane predictions using the hybrid statistical-dynamical framework (GFDL-DecPre; Vecchi *et al.* 2011) and also to apply the hybrid framework to a model system that has shown high multi-year correlations using an alternative approach (UKMO-DePreSys; S10). Additionally, these two models generated a full ensemble of initialized predictions each year, rather than every five years as in many other CMIP5 experiments (Meehl *et al.* 2012), allowing us to more fully explore past performance.

The GFDL decadal climate hindcasts (GFDL-DecPre) are carried out over the period 1961-2011 using the GFDL CM2.1 coupled system (Delworth *et al.* 2006), in which both the atmosphere and the ocean are initialized through a full-field assimilation to bring the state of the coupled model close to observations. The initial conditions are produced with the GFDL fully coupled reanalysis ECDA3.1, which is based on an ensemble Kalman filter (Zhang *et al.* 2007; Zhang and Rosati 2010; Chang *et al.* 2011b) and has been shown to produce a realistic ocean mean state and variability (Chang *et al.* 2012). Ten-member ensembles are produced starting from the first of January every year from 1961 to 2011 and run for ten years. Historical radiative forcing is used for the 1961-2005 period and the Representative Concentration Pathways (RCP) 4.5 scenario for the predictions starting after 2005. A ten-member ensemble of uninitialized runs with the

1 same forcings has also been produced to investigate the impact of initialization. This
2 forecast suite is further discussed in Rosati *et al.* (2012), and its retrospective skill in
3 predicting Atlantic Multidecadal Oscillation-like variability is described in Yang *et al.*
4 (2012).

5 DePreSys (Smith *et al.* 2007) is based on the third Hadley Centre coupled global
6 climate model, HadCM3 (Gordon *et al.* 2000). The UKMO-DePreSys Perturbed Physics
7 Ensemble (PPE; S10) is an updated version that uses a nine-member ensemble of model
8 variants that aims to sample model uncertainties through perturbations to poorly
9 constrained atmospheric and surface parameters. Initial conditions are created by relaxing
10 the model's components toward atmospheric (European Centre for Medium Range
11 Weather Forecasting Analysis and Reanalysis) and oceanic (Smith and Murphy 2007)
12 analysis, with values assimilated as anomalies with respect to the model climate. The
13 purpose of anomaly assimilation is to minimize climate drift after the assimilation is
14 switched off, but this does not totally suppress the bias as discussed in Robson (2011).
15 The ten-year long decadal retrospective forecasts consist of nine-member ensembles
16 starting from the first of November every year from 1960 to 2005. A parallel set of nine
17 uninitialized experiments using the DePreSys-PPE is also used, and is referred to as the
18 UKMO-DePreSys uninitialized forecast runs. The DePreSys experiments do not include
19 future volcanic information in them, only volcanic aerosols from eruptions prior to the
20 initialization; thus, each initial year has a unique suite of uninitialized experiments. We
21 use the UKMO-DePreSys-PPE data, rather than the CMIP5 UKMO-DePreSys output in
22 order to have a comparison to the results of S10.

1 We also perform a two-model average prediction by first running the statistical
2 emulator on the output from each model, and then averaging the predictions of the two
3 models. Previous experience with interannual hurricane forecasts indicates that a two
4 model average can have advantages over each individual model (Vecchi *et al.* 2011).
5 Further work with the full suite of CMIP5 models is underway (Caron *et al.* 2012, in
6 preparation).

7 *C. Lead-dependent climatology:*

8 The statistical hurricane emulator is defined in terms of SST anomalies with respect
9 to the 1982-2005 climatology (Vecchi *et al.* 2011). The initialized and uninitialized
10 model forecasts have their own climatology, which –for initialized forecasts using both
11 models and for uninitialized forecasts using UKMO-DePreSys-PPE – can depend on the
12 lead-time of the forecast. The uninitialized forecasts of DePreSys-PPE have a lead
13 dependent climatology because the history of radiative forcing seen by forecasts
14 verifying on the same year can depend on the initialization year, since no “future”
15 volcanic information is included in these uninitialized experiments. Therefore, we define
16 a different climatology for each experiment (initialized and uninitialized), for each model
17 (GFDL-DecPre and UKMO-DePreSys-PPE). For the initialized model experiments we
18 build a climatology that depends on lead-time by averaging, for each lead-time between
19 one and ten years, the forecasts that verify in the years 1982-2005. We choose this as our
20 reference period for two principal reasons: i) the statistical model of Vecchi et al. (2011)
21 was trained referenced to 1982-2005, and ii) as a trade-off between trying to train over a
22 period in which the observing system used to initialize the forecasts was relatively stable
23 and the desire to have a long record to faithfully define the model drift. Using other

reference periods does not alter the principal results of this manuscript. To compute the model climatology we average all ten ensemble-members for GFDL-DecPre, but since UKMO-DePreSys-PPE is a “perturbed physics ensemble” a different climatology is defined for each of its nine ensemble members. Note that a key impact of subtracting the lead-dependent climatology is to remove a systematic bias that arises in the forecasts as the models drift toward their own mean state when initialized with observations (Stockdale 1997; ICPO 2011). The drift of the models used here is towards each model’s free running climatology, though even after ten years there are regions where the initialized experiments have not yet settled at the free running climatology – these regions tend to roughly coincide with the regions where a potentially predictable decadal signal has been identified in the literature (*e.g.*, Yang *et al.* 2012). A key assumption is that the systematic drift of the models does not depend on initialization period – that is, that the systematic drift does not depend on the changes to the climate observing system that have occurred in the last 50 years. The stationary drift assumption has been shown to be problematic in interannual predictions, where change in observing system can modify the drift, and a suggested solution is to use different lead-dependent climatologies across major changes in observing system (*e.g.*, Kumar *et al.* 2012). The assumption that the drift is stationary will be further discussed in Section IV.

D. Skill measures:

We explore two statistical measures to quantitatively assess retrospective performance: anomaly correlation coefficient (ACC), and mean squared skill score (MSSS). These statistics are not independent, but offer slightly different views of the forecast model skill. ACC is the sample correlation coefficient as a function of lead time

1 t (or an average of lead times), between a set of forecast anomalies F'_j and observed
 2 anomalies O'_j , over $j = 1, \dots, n$ years after removing the mean of each:

$$3 \quad ACC(t) = \frac{\sum_{j=1}^n (F'_j(t) \cdot O'_j(t))}{\sqrt{\sum_{j=1}^n F'^2_j(t) \sum_{j=1}^n O'^2_j(t)}} \quad (Eq.2)$$

4
 5 where $F'_j = F_j - \bar{F}$, $O'_j = O_j - \bar{O}$ and the overbar denotes the time mean over the
 6 climatological period 1982-2005, which is a function of lead time t . ACC values can
 7 range from -1 to 1, and they measure the degree to which large positive and negative
 8 excursions from the mean co-occur in the forecast and verification.

9 The root-mean squared error (RMSE) is often used as a measure of accuracy of the
 10 forecasts. It is defined as the square root of the mean squared error (MSE)

$$11 \quad RMSE(t) = \sqrt{MSE(t)} = \sqrt{\frac{1}{n} \sum_{j=1}^n (F'_j(t) - O'_j(t))^2} \quad (Eq.3)$$

12 We use here a related statistical measure, the mean squared skill score (MSSS;
 13 Murphy 1998) following recommendations by Goddard *et al.* (2012). MSSS is based on
 14 the mean squared error (MSE) between the forecast and the observed climatology and
 15 represents the improvement in accuracy of the forecast over climatology:

$$16 \quad MSSS(t) = 1 - \frac{MSE_F(t)}{MSE_{\bar{X}}(t)} \quad (Eq.4)$$

17 The highest MSSS value of 1 is reached when $MSE_F = 0$ and $MSE_{\bar{X}} \neq 0$.

18 Instead of using climatology as reference forecast one can use the MSE of the
 19 uninitialized projections (MSE_p) to evaluate the improved skill due to initialization:

$$1 \quad MSSS(t) = 1 - \frac{MSE_F(t)}{MSE_P(t)} \quad (Eq.5)$$

2 where a positive MSSS indicates that the initialized forecasts outperform the uninitialized
3 ones. MSSS can be expressed as a function of correlation and conditional bias (Goddard
4 *et al.* 2012), which is useful when interpreting an improvement of skill due to
5 initialization.

6 *E. Assessment of statistical significance:*

7 We explored three different estimates to assess statistical significance of the
8 correlation results against a null of zero correlation, and to compute the confidence
9 intervals of the retrospective correlations. For the estimates of statistical significance the
10 effective number of degrees of freedom (N_{eff}) of the correlation of two time-series (X and
11 Y) was computed using the methodology described in Bretherton *et al.* (1999), using the
12 biased estimates of autocorrelation spectrum of the various time-series:

$$13 \quad N_{eff} = \frac{N}{\sum_{\tau=0}^{N-1} (1-|\tau|/N) r_{\tau}^X r_{\tau}^Y} \quad (Eq.6)$$

14 where N is the number of samples in each time-series, and r_{τ}^X and r_{τ}^Y is the estimate of
15 autocorrelation of each time-series at lag τ . Because of the large autocorrelation of the
16 time-smoothed predicted and observed hurricane time-series at even long lags, the
17 effective degrees of freedom can be considerably smaller than the number of years in the
18 time-series. Typically, when compared to observations, the five-year mean initialized
19 forecasts tend to have between 6-8 effective degrees of freedom and the uninitialized
20 forecasts tend to have between 10-12 effective degrees of freedom – even though there
21 are around fifty years of data that are compared. Without accounting for the strong
22 autocorrelation in these time-series, one would estimate much narrower confidence

intervals and a smaller p -value for the null hypothesis; failure to account for the diminished degrees of freedom can lead to a substantial overestimation of forecast skill.

Though hurricane frequency is not Normally-distributed, we are exploring multi-year averages of hurricane frequency, which allows us to approximate the distribution as Normal. To compute confidence intervals of a correlation we use a two-sided test (since it is possible that initialization could lead to degradation in performance), and use a one-sided test against the null hypothesis of zero correlation (since a significantly negative correlation would be a failure of the forecast system), we have compared the results from three methods:

- i) *Fisher's-z Transformation*: The sample estimate of the correlation coefficient between two time-series (X and Y), $r_{X,Y}$, is transformed using:

$$z_{X,Y} = 0.5 \ln \left[\frac{(1 + r_{X,Y})}{(1 - r_{X,Y})} \right] \quad (\text{Eq.7})$$

The new quantity, $z_{X,Y}$, follows a z distribution with $N_{eff}-3$ degrees of freedom (Fisher 1915, 1924; von Storch and Zwiers, 1999). Using standard z -statistic tables one can estimate the confidence intervals on the mean and test against a null of zero mean from the sample estimate, $z_{X,Y}$. To transform the confidence interval estimates of the z -statistic back to correlation space, we employ the inverse Fisher's- z Transformation:

$$r_{X,Y}^* = \frac{e^{2z_{X,Y}^*} - 1}{e^{2z_{X,Y}^*} + 1} \quad (\text{Eq.8})$$

where $z_{X,Y}^*$ is the estimate of the upper or lower bound on the confidence interval of the z -statistic and $r_{X,Y}^*$ is the estimate of the upper or lower bound on the confidence interval of the correlation coefficient.

ii) *Full distribution of the correlation coefficient:*

Johnson *et al.* (1995) provide the distribution of the sample correlation coefficient R when the population correlation coefficient ρ is equal to zero:

$$p_R(r) = \frac{\Gamma[(n-1)/2]}{\Gamma(1/2)\Gamma[(n-2)/2]} (1-r^2)^{(n-4)/2}, \text{ for } -1 < r < 1 \quad (\text{Eq.9})$$

where $\Gamma(\cdot)$ is the gamma function, n is the sample size. This distribution is symmetric around the zero. By using p_R , we can test the null hypothesis of no correlation at a given significance level α , by checking whether the sample correlation coefficient lies within or outside the rejection or critical region.

iii) *Monte Carlo estimate:* For sample sizes ranging between 2 and 100, we build 100,000 estimates of the distribution of the sample correlation coefficient between two normally-distributed time-series of length N_{eff} and an underlying correlation ρ . We sample underlying correlation coefficients between -1 and 1, at intervals of 0.01. From this Monte Carlo estimate of the probability density function of the sample correlation coefficient, we estimate significance against a null of zero correlation as the probability of a correlation as large as or larger than a particular sample correlation given an underlying correlation of zero. In an analogous manner, we also compute the confidence intervals on the sample correlation given an underlying correlation.

We have compared the three estimates of the confidence intervals on the correlation coefficient and null test against a correlation of zero for the retrospective forecast correlations, and have found that they are consistent with

each other. For simplicity, in the manuscript we only show the estimates from the Fisher's-z transformation.

III. Results

A. Retrospective Hurricane Forecasts:

Figure 1 shows the five-year mean and nine-year mean (centered on the mid-point of each interval) initialized and uninitialized forecasts of North Atlantic hurricane frequency in GFDL-DecPre and UKMO-DePreSys-PPE compared with observations. The observed record of five-year mean hurricane frequency is largely characterized by two distinct states with low values (~5-6 hurricanes per year) in the first half of the record and a shift in the mid-90s (*e.g.*, Elsner et al. 2004, Li and Lund 2012) toward a more active state (~8 hurricanes per year). The uninitialized predictions capture a tendency for an increase in hurricane frequency over the late-20th century, indicating that part of the recent increase in Atlantic hurricane frequency was due to changes in radiative forcing – consistent with other recent findings (*e.g.*, S10; Villarini and Vecchi 2012.b-.c). However, the uninitialized experiments fail to capture the abrupt shift in the mid-1990s. The initialized retrospective forecasts show better qualitative agreement to observations than do the initialized runs, suggesting an improvement from initialization.

Despite the time averaging, both observations and the model predictions have year-to-year variability in five-year North Atlantic hurricane frequency, which complicates detection of decadal changes (Figure 1). The year-to-year variations in the multi-year initialized forecasts are larger than that in observations, even though the forecasts are ensemble averages. This result is particularly striking given that the statistical emulator should only recover a fraction of the observed variance, and suggests that the initialized

1 forecasts have too much internal variability. An alternative interpretation, which is
2 discussed further in Section III.C below, is that the initial conditions for each year's
3 initialization are persisted too strongly, so that each initialization year's climate reflects
4 on the average of multiple subsequent years.

5 The anomaly correlation between the observed hurricane counts and the models
6 predictions for both initialized and uninitialized experiments is shown in Figure 2 for
7 five-year and nine-year means. A persistence forecast is given as a reference test forecast,
8 where the five-year (nine-year) mean persistence is defined as the observed average over
9 the five (nine) years that precede the model's initialization (persisting the SSTA indices
10 does not improve the performance of the persistence null model, with correlations
11 ranging between 0.16-0.4 depending on the SST dataset used). So, for example, the
12 persistence forecast for the lead 2-6 forecast centered in 1992 (*e.g.*, initialized in 1989) is
13 the observed hurricane count averaged over 1984-1988. Consistent with Figure 1, at lead
14 2-6 the initialized retrospective predictions show higher correlations than the uninitialized
15 ones, for both models. The values are significantly different from zero and exceed the
16 values given by persistence, which is not the case for the uninitialized predictions.
17 Comparable skill is found between the two models, slightly higher in UKMO-DePreSys;
18 these retrospective correlations are comparable to those reported in S10 using an
19 alternative methodology applied to DePreSys-PPE. Computing the two-model mean
20 increases the signal-to-noise ratio, leading to higher correlations than in either individual
21 model. At lead 2-10, all the predictions outperform the persistence forecast. The decadal
22 correlations are nominally higher in the initialized retrospective predictions than in the
23 uninitialized, with the largest values, exceeding 0.8, when taking the two-model mean.

1 This decadal skill does not come only from the first few years since the correlations at
2 lead 6 to 10 are also large (Figure 2), although the improvement due to initialization is
3 not as clear. At lead 6-10, GFDL-DecPre shows larger correlations for the initialized
4 predictions but UKMO-DePreSys indicates higher values for the uninitialized runs,
5 yielding undistinguishable values between the initialized and the non-initialized
6 experiments for the two-model mean.

7 These results suggest that coupled GCMs that account for both changes in initial state
8 and radiative forcings can lead to skillful multi-year retrospective predictions of
9 hurricane frequency. The nominal improvement due to initialization should, however, be
10 interpreted with care given the large confidence intervals associated with the point
11 estimates of the correlations (Figure 2). As discussed above in Section II.E, although the
12 observed record is 50-years long, because of the large autocorrelation of the time series
13 each year is not independent from those nearby. Hence, the effective number of degrees
14 of freedom is largely reduced to less than ten for most lead times, as indicated on Figure
15 2, based on Bretherton *et al.* (1999). Therefore, even if the initialized predictions give a
16 correlation that is statistically different from climatology and is nominally higher than in
17 the uninitialized predictions, the large confidence intervals indicate that the retrospective
18 correlation of the initialized forecasts is not different from persistence or the uninitialized
19 experiments at $p=0.1$. Some of the correlations of the initialized forecasts are
20 significantly larger than the non-initialized experiments at $p=0.2$.

21 The non-significance of the difference between the initialized and non-initialized
22 correlations does not depend strongly on the effective sample size, as long as some level
23 of autocorrelation is assumed. We recomputed the confidence intervals on the sample

1 correlations using an unrealistic assumption that two years were needed for each new
2 degree of freedom, and the initialized to uninitialized correlation differences were still
3 not significantly different at $p=0.1$. If we assume, even more unrealistically, that a new
4 degree of freedom is achieved every 1.5 years, then the differences between the
5 initialized and uninitialized experiments are significant at $p=0.1$. However, we wish to
6 stress that these perturbation experiments yield an extremely unrealistically high estimate
7 of the number of degrees of freedom, considering we are exploring five-year running
8 averages of quantities with a pronounced trend and interdecadal variation. The record is
9 too short, and the difference between initialized and uninitialized correlations too small,
10 to yield a statistically significant difference.

11 Improvement from initialization on the two-model mean lead 2-6 forecast is close to
12 being significant even at $p=0.1$, suggesting potentially higher confidence in multi-model
13 ensembles. For the lead 2-6 and 2-10 forecasts, for both model systems there is a
14 consistent nominal improvement of retrospective correlation from initialization relative to
15 the uninitialized experiments. Because of this, and because of the small sample size, we
16 speculate that the lack of significance at $p=0.1$ may reflect a “lack of power” by the
17 significance test, rather than a “lack of effect” from initializing (Johnson 1999). For the
18 lead 6-10 forecast, however, the nominal difference between the initialized and non-
19 initialized forecasts changes sign (there is nominal indication of improvement in GFDL-
20 DecPre, but a nominal degradation in UKMO-DePreSys-PPE), so we interpret the lack of
21 significance in this case as indicating a lack of effect from initialization. Therefore, it
22 appears that the nominal improvement in the lead 2-10 forecast arises in the first part of
23 the decade, and represents potential multi-year forecast skill rather than decadal skill.

1 A lagged ensemble approach, in which past forecasts are used to augment the
2 effective ensemble size of more recent forecasts (*e.g.*, by creating a forecast where the
3 current year's lead 1-5 and the previous year's lead 2-6 forecasts are averaged), can lead
4 to increase in forecast performance (*e.g.*, Vecchi *et al.* (2011) showed improvement in
5 interannual hurricane forecasts from lagged ensembles). We explored the impact of
6 lagged ensembles in the retrospective hurricane forecasts (not shown) at lags of up to
7 three years (*i.e.*, averaging lead 1-5, 2-6 and 3-7 verifying the same years together)
8 resulted in nominal improvements in the correlation coefficient (on the order of 0.02-
9 0.05). However, the smoothing induced by the lagged ensemble led to a further reduction
10 of degrees of freedom. Since the uncertainty in a correlation estimate increases with
11 decreasing correlation or sample size, the uncertainty estimates on the correlation
12 coefficient did not show substantial change: even after lagged-ensemble averaging the
13 retrospective correlation of the uninitialized and initialized forecasts were in each other's
14 confidence intervals.

15 As a complement to the skill estimate using ACC, we show in Figure 3 the MSSS for
16 various five-year mean and nine-year mean leads. Both the improvement relative to
17 climatology (Eq.4) and that due to initialization (Eq.5) are indicated on the x- and y-axis,
18 respectively. None of the retrospective initialized forecasts has a negative MSSS on the
19 x-axis, which indicates at least a nominal improvement relative to climatology. An
20 improvement due to initialization is also suggested at all leads in GFDL-DecPre, and at
21 most leads except 5-9 and 6-10 in UKMO-DePreSys, leading to a smaller MSSS at those
22 lead times for the two-model mean. Both models indicate an improved skill at decadal
23 scale due to initialization, with the highest values in UKMO-DePreSys. As shown in

1 Goddard *et al.* (2012), the MSSS is a function of both the correlation and the conditional
2 bias, and the higher MSSS due to initialization is mainly due to a reduction of the
3 conditional bias that is large in the uninitialized predictions.

4 *B) SST-source of hurricane forecast skill*

5 Our hurricane frequency index is based on SST averaged over the tropical Atlantic
6 and over the global tropics (Eq.1), so both quantities are potential sources for the better
7 predictability in the initialized forecasts. We can explore retrospective forecasts and skill
8 measures of these two indices with hope of finding the role each had in recovering the
9 past history of hurricane activity (Figure 4). Overall, there is no indication that
10 retrospective forecasts of tropical-mean SST are improved by initializing the coupled
11 GCMs (upper panels, Figure 4), with the relatively monotonic warming of the tropics
12 dominating the observed and modeled signals. The dominance of the long-term trend in
13 both SST indices cuts the effective degrees of freedom severely, to the point where for
14 tropical-mean SST interpretation of correlation as a skill metric is likely too ambiguous
15 to be useful. The GFDL-DecPre system has marginally higher retrospective correlation in
16 both SST indices than does UKMO-DePreSys, likely due to inclusion of future volcanic
17 information in its radiative forcing (Table 1). However, this nominally larger skill in
18 GFDL-DecPre for the two SST indices does not translate into even nominal increase of
19 the hurricane forecasts (Figure 2) since the volcanic signals are primarily spatially
20 uniform. Across both model systems there is a consistent nominal improvement of
21 retrospective correlation of Atlantic MDR SST predictions from initialization, but the
22 effect is small relative to the number of degrees of freedom. Only in the GFDL-DecPre
23 does the initialized forecast of MDR SST approach a significant improvement over a

1 persistence forecast. Because of the dominance of a quasi-monotonic trend, for tropical-
2 mean SST all the forecast methods (initialized and uninitialized GCM forecasts and
3 persistence) yield comparable results. For both SST indices all of the forecast
4 methodologies lead to statistically significant retrospective correlations against a null of
5 zero correlation, again largely because of the dominance of a trend.

6 The results in Figure 4 suggest that the nominal improvement in retrospective
7 correlation from initialization came from improvements to forecast of Atlantic MDR
8 SST. However, since the time series of each SST index includes a substantial component
9 that is coherent across both indices, and since the hurricane frequency emulator is based
10 on the difference between the two indices, interpreting the source of hurricane
11 predictability from each index is not necessarily straightforward, as was noted in Vecchi
12 *et al.* (2011). An alternative approach to assessing influence of each index on the role of
13 initialization on forecast skill is to use values of one index from the initialized
14 experiments and the other from the uninitialized experiments. For example, taking values
15 for SST_{MDR} from the initialized experiment, but keeping the SST_{TROP} from the
16 uninitialized one, yields comparable hurricane retrospective forecast results (Figure 5a) to
17 when both indices are taken from the initialized experiments (Figure 2). The impact of
18 initialization on SST_{MDR} yields five-year mean fluctuations of this hurricane frequency
19 index that show rather good agreement with observations for both models with a
20 correlation of 0.70 and 0.59 in GFDL-DecPre and UKMO-DePreSys, respectively (both
21 significantly different from zero correlation at $p < 0.05$) at lead 2-6. Using values for
22 SST_{MDR} from the uninitialized experiments but those of SST_{TROP} from the initialized
23 experiments leads to very different results (Fig 5.b). The correlation drops to 0.21 in

1 GFDL-DecPre and to 0.43 in UKMO-DePreSys, with neither correlation significantly
2 different from $\rho=0$ (even at $p<0.2$) nor either model able to reproduce the observed sharp
3 increase in the mid 90s. This indicates that the nominal improvement in correlation in the
4 initialized multi-year predictions results from a better representation of the Atlantic main
5 development region when initializing the coupled models, with little beneficial impact
6 from initialized predictions of the global mean tropical SST.

7 For the GFDL-DecPre system the difference in retrospective correlation when
8 swapping initialized/uninitialized SST_{MDR} and SST_{TROP} is significant at $p<0.1$. Note in
9 Figure 5.b there is a large increase in hurricane frequency around 2005 in GFDL-DecPre,
10 as appeared in Fig1.a. This increase, which we currently consider to be spurious, is a
11 large contributor to the reduction in correlation from the impact of initialization on
12 tropical-mean SST in the GFDL model. There is a coincidence between the global
13 implementation of the “Array for Real time Geostrophic Oceanography” (or Argo)
14 drifting float profiles in 2003 and the spurious shift of nine-year forecasts centered
15 around 2005-2006, suggesting that enhanced observational sampling after 2003 may have
16 led to a change in the lead-dependent climatology. Experiments are underway to test this
17 possibility. The lack of such a spurious increase in UKMO-DePreSys could arise from
18 different initialization processes, or from the fact that the last initialized forecast in
19 UKMO-DePreSys begins in 2006 – so the late spike would not be evident. Were the
20 introduction of Argo found to be the driver of this spurious increase, in addition to
21 developing methods to minimize the impact of observing system changes, the impact of
22 other large changes to the observing system must also be explored (*e.g.*, the introduction
23 of altimetry in the early 1990s and the completion of the TAO array in the mid-1990s).

1 *C) Role of the mid-1990s climate shift:*

2 The nominal improvement in skill due to initialization should be interpreted with
3 care. Even if the initialized retrospective predictions outperform climatology at almost all
4 lead times (Figure 3), the skill could still come from persistence – just persistence that
5 cannot be captured with our observationally-based persistence model. Figure 6a and 6b
6 compare the retrospective predictions of hurricane frequency for five-year means ranging
7 between lead 1-6 to lead 6-10. The forecasts at each lead show a tendency to have a
8 systematic one year shift with respect to the preceding lead, with the mid-1990s shift in
9 each model trailing in time for longer leads rather than capturing the observed 1995 shift
10 (*e.g.*, Elsner et al. 2004, Li and Lund 2012) at the right time. By performing change point
11 analysis (Pettitt test) on the models' retrospective predictions, we find a shift in forecasts
12 initialized in 1991 in UKMO-DePreSys and forecasts initialized in 1995 in GFDL-
13 DecPre. This tendency for forecasts to lock across the shift can be seen more clearly
14 when the same time series are plotted as a function of initialization year instead of
15 verification time (Fig 6c and 6d): forecasts initialized the same year are very similar to
16 each other, independent of when they verify. Notice that the mid-90s shift for each model
17 appears at the same initialization year for all lead times, as does the potentially spurious
18 mid-2000s shift in GFDL-DecPre.

19 Up to now we have been largely comparing the results of forecasts initialized
20 different years at the same lead, without focusing on the evolution of hurricane counts of
21 each forecast as the lead increases. A correct forecast of the mid-1990s climate shift
22 would have indicated at some point prior to the shift that there was an increased
23 probability of hurricane frequency increasing in time. For example, if a forecast

1 initialized in early 1991 showed counts averaged in 1992-1996 that were larger than
2 those in 1991, or an increased number of ensemble members with large increases, one
3 would have evidence for a future shift. Do these two forecast systems produce such a
4 shift? Figure 7 shows that in the observational record, reflecting the rapid increase in
5 frequency in 1995, the difference in hurricane counts averaged over the five years
6 following the years 1991 through 1994 exceeded the counts over each of those years by
7 an unusually large amount, relative to the distribution over the 1961-2006 period.
8 However, neither forecast system (colored lines in Figure 7) shows a tendency for their
9 forecasts to increase in time relative to the first forecast year when initialized in the early
10 1990s. In fact, there is a nominal tendency for these forecasts to decrease in time from the
11 first forecast year, relative to the distribution of tendencies across all initialization dates,
12 1961-2006. That is, the models did not forecast a *tendency* towards higher frequency in
13 the mid-1990s (Figure 7), even though the sequence of forecast *values* exhibits a climate
14 shift in the mid-1990s (Figures 1, 6).

15 To further highlight the influence of the mid-1990s shift on the retrospective skill
16 estimation, we explore forecast performance after removing the mid-90s shift from both
17 the forecasts and the observations. The shift is “removed” by simply referencing each
18 period before and after the 1994-1995 shift to its own climatology; for instance, the time-
19 mean hurricane count preceding 1995 is removed from all years before 1995, and the
20 time-mean hurricane count following 1995 is removed from all years after 1995. We note
21 that using each model’s change-point instead of 1995 does not affect the character of the
22 results. Figures 8 and 9 indicate that removing the shift leads to a substantial reduction of
23 correlation in the initialized predictions at lead 2-6 (particularly for UKMO-DePreSys-

1 PPE), and no indication of skill beyond that lead time, further confirming that the decadal
2 signal is dominated by the trend that arises from the existence of the mid-90s change
3 point. Therefore, future real (as opposed to the retrospective forecasts explored here)
4 multi-year and decadal predictions of hurricane frequency should not be expected to
5 show the same skill as over the 1961-2011 period unless there are change points of
6 similar character to the mid-1990s shift. Our results are encouraging for the feasibility of
7 multi-year forecasts of hurricane frequency with the current prediction systems.
8 However, this analysis highlights that substantial challenges remain – or, viewed more
9 optimistically, that it is possible to improve the performance of the system beyond its
10 current capability.

11 An interesting side effect of removing the mid-1990s shift is to increase the effective
12 degrees of freedom, narrowing the confidence intervals associated with the point
13 estimates of the correlation coefficient (compare Figures 2 and 9). In addition, the
14 retrospective correlation in the uninitialized forecasts without change-point disappeared –
15 since it largely arose from the projection of the observed shift onto the models' forced
16 trend over this period. In this modified context, there is now indication that for the GFDL
17 model and the two-model ensemble the correlations (although lower than in the case
18 including the shift; Figure 2) are significantly higher than those of the uninitialized
19 versions of the model at lead 2-6. That is, there is significant (at $p < 0.1$) indication that
20 GFDL-DecPre and the two-model ensemble may be able to predict the types of variations
21 in hurricane frequency that occurred in the early-1980s and early-1990s better than the
22 uninitialized experiments. In Figure 2, the nominal improvement from initialization in the
23 correlation of the lead 2-6 and lead 6-10 mean hurricane counts in GFDL-CM2.1 was

1 larger than that for the lead 2-10 forecasts; this may reflect the ability of GFDL-CM2.1 to
2 retrospectively forecast some multi-year variations beyond the 1994-1995 climate shift –
3 which is the dominant signal in the nine-year running counts. This further highlights the
4 limitations of a data record that is short relative to the dominant timescales in order to
5 assess the impact of multi-year forecast skill. While it is entirely possible that some of the
6 non-significant differences between the initialized and uninitialized models shown in
7 Figures 2 and 3 could become significant from a longer record, it is also possible that the
8 impact of initialization could also decrease and remain non-significant in a longer record.

10 **IV Summary and Discussion**

11 Predictions of North Atlantic hurricane frequency were investigated in two global
12 coupled models initialized towards estimates of the observed climate state. We find
13 statistically significant retrospective correlation of multi-year to decadal initialized
14 hurricane frequency forecasts by accounting for both initialization and radiative forcing
15 changes. The two systems explored, GFDL-DecPre and UKMO-DePreSys-PPE, show
16 comparable skill. The two-model mean had the best skill, encouraging the pursuit of
17 broader multi-model studies (*e.g.*, Caron *et al.* 2012); lagged averages lead to nominal
18 correlation increases. The retrospective correlations from initialized multi-year hurricane
19 forecasts are comparable to those reported in Smith *et al.* (2010; S10) using an alternative
20 methodology.

21 Taken together, our results and those of S10 indicate that initializing a climate model
22 and accounting for radiative forcing changes, together, can lead to significant
23 retrospective skill in multi-year initialized (relative to climatological forecasts). The

1 performance of the initialized forecasts was nominally better than that of uninitialized
2 forecasts, both in correlation and in MSSS (Goddard *et al.* 2012). However, because of
3 the short observational record and the persistent character of the time series, the
4 confidence intervals associated with all the forecasts are large, and the difference
5 between initialized and uninitialized forecasts is not statistically significant at $p=0.1$
6 (although some are at $p=0.2$). Because of the consistency of correlations across studies
7 and the visual improvement, we hypothesize that lack of significant improvement from
8 initialization may indicate of lack of “power” (*i.e.*, the probability that the test will
9 correctly reject the null hypothesis) by the statistical test (arising from too few degrees of
10 freedom and a relatively strong correlation arising from radiative forcing alone) rather
11 than a lack of effect of initialization (*e.g.*, Johnson 1999). Additional years could lead to
12 enhancement of our confidence; however, the large autocorrelation of the time series
13 indicates that we require about seven years of data to gain a degree of freedom – so many
14 years will be required to improve our confidence, even if we include the past 50 years in
15 future estimates of forecast skill.

16 The observed time series of North Atlantic hurricane frequency is dominated by a
17 strong and abrupt rise in 1995 leading to a trend over the 1961-2011 period. The high
18 correlations of the retrospective predictions of North Atlantic hurricane frequency depend
19 on the presence of this shift. While predictions from both models are for more hurricanes
20 after the mid-90s than before, the increase is not actually predicted by the evolution of the
21 models, but is present in the initial state (*i.e.*, forecasts initialized after the shift exhibited
22 by each model remain high, but those initialized prior do not show the shift; Fig. 6-7).
23 That is, the large retrospective skill estimates (Figures 2-3) do not come from predicting

1 the dynamical evolution of the climate system resulting in the hurricane frequency shift,
2 but from “recognizing” that a climate shift has occurred and persisting that shift. This
3 behavior mirrors experience in seasonal forecasts of El Niño, where transition from
4 climatological conditions to a warm ENSO state can be problematic to predict (*e.g.*,
5 Landsea and Knaff 2000; Vecchi *et al.* 2006), and successful forecasts often reflect the
6 continued updating of subsurface conditions. This reduces our confidence that the onset
7 of a similar shift in a near future could be successfully predicted with current prediction
8 systems. It also highlights the need to better understand the origin of the change point in
9 the observations and assess whether the modeled mechanisms are consistent with those in
10 the real world (*e.g.*, Robson *et al.* 2012).

11 Despite high correlation values, the mean retrospective skill of these forecasts may
12 provide a poor and even misleading guide to the future performance. In the absence of a
13 major climate shift, like the 1994-1995 shift, the long-term estimates of correlation (*e.g.*,
14 0.6-0.9) are not representative, and the lower retrospective correlations assessed after
15 removing the shift (*e.g.*, 0-0.4; Figs. 8-9) may be closer to those one should expect.

16 Neither model system successfully predicts that the highest values of observed five-
17 year hurricane frequency that appear in the mid-2000s. GFDL-DecPre shows a
18 comparable rise but five to ten years later than observed, whereas UKMO-DePreSys
19 shows a more modest increase with a several-year delay as well. Forecasts with GFDL-
20 DecPre that extend past the present suggest an increase in hurricane frequency through
21 the mid-2010s (Fig. 1). However, observations have been tending in the opposite
22 direction, with recent years being less active than those in the mid-2000s. This period
23 coincides with a fundamental change in the ocean observing system, with the global

1 introduction of Argo floats after 2003 bringing a considerably better coverage of the
2 surface and subsurface ocean. Changes in observing systems have previously impacted
3 the behavior of initialized forecasts, in part by changing the character of the initialized
4 model's drift (*e.g.*, Kumar *et al.* 2012); therefore the introduction of Argo could impact
5 the lead-dependent climatology.

6 Thus, we hypothesize that this increase predicted by with GFDL-DecPre is spurious,
7 and reflects the impact of Argo data on the GFDL-DecPre drift. To test this hypothesis a
8 set of experiments was performed in which Argo data was withheld from the
9 initialization scheme of GFDL-DecPre after 2004. The predicted abrupt increase after
10 2004 is severely reduced when Argo is removed (Fig. 10), largely because of changes to
11 model drift in regions that were poorly observed prior to Argo. These experiments
12 support our hypothesis, so a more plausible prediction for the coming years is that shown
13 in the left panel of Figure 5, in which there is a tendency for relative stability to a
14 reduction of hurricane frequency in coming years. Changes in drift (lead-dependent
15 climatology) arising from the introduction of Argo impact the character of predictions of
16 tropical-mean and global-mean temperature in the GFDL-DecPre system, leading to
17 spuriously cold predictions of both if a single lead-dependent climatology is used to
18 analyze the pre- and post-Argo period. We speculate that related errors may arise in this
19 other prediction systems due to observing system changes. Methodologies to deal with
20 the impact of observing system changes on drift must be developed in order to fully
21 realize the potential of multi-year predictions; as the post-Argo record lengthens,
22 motivated by Kumar *et al.* (2012), a potential solution is to use different lead-dependent
23 climatologies for the pre- and post-Argo period. In addition, the impact of other

1 observing system changes bear exploration, such as the introduction of the Pacific
2 Tropical Atmosphere-Ocean moored buoy array in the early-1990s (McPhaden 1993) and
3 expendable bathythermographs in the late 1960s. Interpretation of forecasts needs to be
4 keenly constrained by our knowledge of changing observing practices both in the
5 predictands (*e.g.*, Vecchi and Knutson 2008, 2011; Landsea *et al.* 2010; Villarini *et al.*
6 2011b) and in the observations used to initialize the climate model (*e.g.*, Zhang *et al.*
7 2007; Kumar *et al.* 2012).

8 Identifying the source of skill in retrospective predictions is key to the success of
9 future forecasts. Recent studies (Mann and Emanuel 2007; Evan *et al.* 2009; S10;
10 Villarini and Vecchi 2012b,c) have argued that the recent (since the 1980s) increase of
11 Atlantic hurricane activity was not caused by internal variability alone but also included
12 an externally-forced component driven largely by changing aerosol concentrations. Our
13 results partially support this interpretation, indicating high correlations (significantly lead
14 2-10) in the uninitialized forecasts. Yet the sharp mid-90s increase in Atlantic hurricane
15 frequency is not retrospectively predicted in the uninitialized experiments. Its better
16 representation in the initialized predictions could be interpreted as an indication of a key
17 role for internal variability in the mid-1990s shift, supporting various studies (*e.g.*, Zhang
18 and Delworth 2005,2006,2009; Robson *et al.* 2012; Yeager *et al.* 2012; Msadek *et al.*
19 2012). However, the nominal improvement from initialization could also reflect a failure
20 in the radiative forcing/response in these models that is corrected when they are
21 constrained with observations.

22 Our results indicate that the impact of initialization on forecasts of the Atlantic
23 main development region (MDR) relative to the tropics was key to the higher skill in the

1 initialized forecasts (Figures 4 and 5). Zhang and Delworth (2006) suggested that multi-
2 year changes in hurricane activity could be driven by changes to the heat-transport over
3 the entire North Atlantic. S10 and Dunstone *et al.* (2011) further suggested that the
4 subpolar North Atlantic was the main source of multi-year predictability of Atlantic
5 hurricane frequency. The North Atlantic also stands out as the region where initialized
6 forecasts outperform uninitialized ones in the GFDL model (Rosati *et al.* 2012; Yang *et*
7 *al.* 2012; Msadek *et al.* 2012), suggesting a potential link between North Atlantic
8 variability and Atlantic hurricane predictability in GFDL DecPre. Further, Kang *et al.*
9 (2008) showed that changes in the North Atlantic could lead to changes in atmospheric
10 circulation over the tropical Atlantic in GFDL CM2.1. However, in our retrospective
11 forecasts of hurricane activity, the relevant source of skill must have been present in
12 tropical Atlantic SST – so any role for extratropical forcing must involve a subsequent
13 change to tropical Atlantic SST. Thus, improved representation of processes controlling
14 tropical Atlantic climate (*e.g.*, Doi *et al.* 2012) are key to enhanced skill in forecasts of
15 hurricane activity by systems like those used here.

16
17 **Acknowledgments:**

18 We are grateful to Doug Smith (UK Met Office) for making the UKMO-DePreSys PPE
19 data available. We thank Ming Zhao and Tom Knutson for comments and suggestions.

20
21 **References:**

1 Alessandri, A., A. Borrelli, S. Gualdi, E. Scoccimarro, and S. Masina, Tropical cyclone
 2 count forecasting using a dynamical seasonal prediction system: Sensitivity to
 3 improved ocean initialization, *Journal of Climate*, **24**, 2963-2982, 2011.

4 Bender, M.A., T.R. Knutson, R.E. Tuleya, J.J. Sirutis, G.A. Vecchi, S.T. Garner, and I.M.
 5 Held, 2010: Model impact of anthropogenic warming on the frequency of intense
 6 Atlantic hurricanes. *Science* **327**, 454–458.

7 Booth, B.B., N.J. Dunstone, P.R. Halloran, T. Andrews, and N. Bellouin, 2012: Aerosols
 8 implicated as a prime driver of twentieth-century North Atlantic climate variability.
 9 *Nature*, **484**, 228-232.

10 Bretherton, C.S., M. Widmann, V.P. Dymnikov, J.M. Wallace, and I. Bladé, 1999: The
 11 Effective number of spatial degrees of freedom of a time-varying field. *J.*
 12 *Climate*, **12**, 1990-2009.

13 Broccoli, A.J., and S. Manabe, 1990: Can existing climate models be used to study
 14 anthropogenic changes in tropical cyclone climate? *Geophys. Res. Lett.*, **17**, 1917-
 15 1920.

16 Camargo, S.J., A.G. Barnston, P. Klotzbach, and C.W. Landsea, 2007a: *Seasonal tropical*
 17 *cyclone forecasts*, World Meteorological Organization Bulletin, 56, 297-309.

18 ———, K.A. Emanuel, and A.H. Sobel, 2007b: Use of a genesis potential index to
 19 diagnose ENSO effects on tropical cyclone genesis. *J. Climate*, **20**, 4819–4834.

20 ———, M. Ting, and Y. Kushnir, 2012: Influence of local and remote SST on North
 21 Atlantic tropical cyclone potential intensity. *Climate Dynamics (submitted)*.

22 Caron, J.-P., and coauthors: Multi-year hurricane forecasts using the CMIP5 ensemble. In
 23 preparation.

1 Chang, C.-Y., J.C.H. Chiang, M.F. Wehner, A. Friedman, and R. Ruedy, 2011a: Sulfate
2 aerosol control of tropical Atlantic climate over the 20th century. *Journal of Climate*,
3 **24**, 2540–2555.

4 Chang, Y.-S., S. Zhang, and A. Rosati, 2011b: Improvement of salinity representation in
5 an ensemble coupled data assimilation system using pseudo salinity profiles.
6 *Geophysical Research Letters*, **38**, L13609, DOI:10.1029/2011GL048064.

7 Chang, Y.-S., S. Zhang, A. Rosati, T. Delworth, and W. F. Stern, 2012: An assessment of
8 oceanic variability for 1960–2010 from the GFDL ensemble coupled data
9 assimilation, *Climate Dynamic* (in press).

10 Chen, J.H., and S.J. Lin, 2011: The remarkable predictability of inter-annual variability
11 of Atlantic hurricanes during the past decade. *Geophysical Research Letters*, **38**
12 (L11804), doi:10.1029/2011GL047629.

13 Chikamoto Y., M. Kimoto, M. Ishii, T. Mochizuki, T. T. Sakamoto, H. Tatebe, Y.
14 Komuro, M. Watanabe, T. Nozawa, H. Shiogama, M. Mori, S. Yasunaka, and Y.
15 Imada, 2012: An overview of decadal climate predictability in a multi-model
16 ensemble by climate model MIROC. *Clim. Dyn.* doi:10.1007/s00382-012-1351-y.

17 Collins, M., et al., 2006: Interannual to decadal climate predictability in the North
18 Atlantic: A multimodel-ensemble study, *J. Climate*, **19**, 1195–1203.

19 Delworth, T. L., and Coauthors, 2006: GFDL's CM2 global coupled climate models. Part
20 I: Formulation and simulation characteristics, *Journal of Climate*, **19**, 643–674.

21 ———, and Dixon K.W., 2006: Have anthropogenic aerosols delayed a greenhouse gas-
22 induced weakening of the North Atlantic thermohaline circulation? *Geophysical*
23 *Research Letters*, **33**, L02606, DOI:10.1029/2005GL024980.

1 Doi, T., G.A. Vecchi, A.J. Rosati and T.L. Delworth, 2012: Tropical Atlantic biases in
2 the mean state, seasonal cycle, and interannual variations for a coarse and high
3 resolution coupled climate model. *J. Climate*, doi:10.1175/JCLI-D-11-00360.1

4 Dunstone, N. J., D. M. Smith, and R. Eade, 2011: Multi-year predictability of the tropical
5 Atlantic atmosphere driven by the high latitude North Atlantic Ocean. *Geophys.*
6 *Res. Lett.*, **38**, L14701, doi:10.1029/2011GL047949.

7 Elsner, J.B., and T.H. Jagger, 2006: Prediction models for annual U.S. hurricane counts,
8 *Journal of Climate*, **19**, 2935-2952.

9 ———, X. Niu, and T.H. Jagger, 2004: Detecting shifts in hurricane rates using a Markov
10 Chain Monte Carlo approach, *Journal of Climate*, **17**, 2652–2666.

11 Emanuel, K. A., 1987: The dependence of hurricane intensity on climate. *Nature* **326**,
12 483–485.

13 ———, Increasing destructiveness of tropical cyclones over the past 30 years. *Nature*, **436**,
14 686–688, 2005.

15 ———, 2007: Environmental factors affecting tropical cyclone power dissipation. *J. Clim.*
16 **20**, 5497–5509.

17 ———, R. Sundararajan, and J. Williams, Hurricanes and global warming—Results from
18 downscaling IPCC AR4 simulations. *Bull. Amer. Meteor. Soc.*, **89**, 347–367,
19 2008.

20 Evan, A.T., D.J. Vimont, A.K. Heidinger, J.P. Kossin, and R. Bennartz, 2009: The role of
21 aerosols in the evolution of tropical North Atlantic Ocean temperature anomalies.
22 *Science*, **324**, 778–781.

1 Fisher, R.A., 1915: Frequency distribution of the values of the correlation coefficient in
2 samples from an indefinitely large population. *Biometrika*, **10**, 507-521.

3 ———, 1924: The distribution of the partial correlation coefficient. *Metron*, **3**, 329-332.

4 Goddard L., A. Kumar, A. Solomon, D. Smith, G. Boer, P. Gonzalez, V. Kharin, W.
5 Merryfield, C. Deser, S. Mason, B. Kirtman, R. Msadek, R. Sutton, E. Hawkins,
6 T. Fricker, G. Hegerl, C. Ferro, D. Stephenson, G.A. Meehl, T. Stockdale, R.
7 Burgman, A. Greene, Y. Kushnir, M. Newman, J. Carton, I. Fukumori, T.
8 Delworth, 2012: A verification framework for interannual-to-decadal predictions
9 experiments, *Climate Dynamics*, under revision

10 Gordon, C., C. Cooper, C. Senior, H. Banks, J. Gregory, T. Johns, J. Mitchell, and R.
11 Wood, 2000: The simulation of SST, sea ice extents and ocean heat transports in a
12 version of the Hadley Centre coupled model without flux adjustments, *Climate*
13 *Dynamics*, **16**, 147–168.

14 Gray, W.M., 1984: Atlantic seasonal hurricane frequency. Part I: El Niño and 30 mb
15 quasi-biennial oscillation influences, *Monthly Weather Review*, **112**, 1649-1668.

16 Griffies, S.M., and K. Bryan, 1997a: Predictability of North Atlantic multidecadal
17 climate variability, *Science*, **275**(5297), 181.

18 ———, and K. Bryan, 1997b: A predictability study of simulated North Atlantic
19 multidecadal variability, *Climate Dynamics*, **13**, 459–487

20 Gualdi, S., E. Scoccimarro, and A. Navarra, 2008: Changes in tropical cyclone activity
21 due to global warming: Results from a high-resolution coupled general circulation
22 model. *J. Climate*, **21**, 5204–5228.

1 Hawkins, E., and R. Sutton, 2009: The potential to narrow uncertainty in regional climate
2 predictions. *Bulletin of the American Meteorological Society*, **90**, 1095–1107.

3 ICPO (International CLIVAR Project Office), 2011: Decadal and bias correction for
4 decadal climate predictions. January. International CLIVAR Project Office,
5 CLIVAR Publication Series No.150, 6pp. Available from
6 http://eprints.soton.ac.uk/171975/1/150_Bias_Correction.pdf

7 Jarvinen, B.R., C.J. Neumann, and M.A.S. Davis, 1984: A tropical cyclone data tape for
8 the North Atlantic Basin, 1886–1983: Contents, limitations, and uses. Tech.
9 Memo. NWS NHC 22, National Oceanic and Atmospheric Administration, 24 pp.

10 Johnson, D.H., 1999: The insignificance of significance testing. *J. of Wildlife*
11 *Management*, **63**(3), 763-772.

12 Johnson, N.L., S. Kotz, and N. Balakrishnan, 1995: *Continuous Univariate Distributions*
13 (volume 2), Wiley, 752 pages.

14 Kalnay and coauthors, The NCEP/NCAR 40-year reanalysis project. *Bull. Amer.*
15 *Meteorol. Soc.*, 77(3), 437-471, 1996.

16 Kim, H.-M., and P.J. Webster, Extended-range seasonal hurricane forecasts for the North
17 Atlantic with a hybrid dynamical-statistical model, *Geophys. Res. Lett.*, **37**,
18 L21705, doi:10.1029/2010GL044792, 2010.

19 Klotzbach, P.J., and W.M. Gray, 2009: Twenty-five years of Atlantic basin seasonal
20 hurricane forecasts, *Geophysical Research Letters*, **36** (L09711),
21 doi:10.1029/2009GL037580.

1 Knight, J.R., R.J. Allan, C.K. Folland, M. Vellinga, and M.E. Mann, 2005: A signature of
2 persistent natural thermohaline circulation cycles in observed climate. *Geophys.*
3 *Res. Lett.*, **32**, L20708, doi:10.1029/2005GL024233.

4 Knutson, T.R., J.J. Sirutis, S.T. Garner, I. Held, and R.E. Tuleya, 2007: Simulation of
5 recent increase of Atlantic hurricane activity using an 18-km-grid regional model.
6 *Bull. Amer. Meteor. Soc.*, **88**, 1549–1565.

7 ———, ———, ———, G.A. Vecchi, and I. Held, 2008: Simulated reduction in Atlantic
8 hurricane frequency under twenty-first- century warming conditions. *Nat. Geosci.*,
9 **1**(6), 359–364.

10 ———, *et al.*, 2010: Tropical cyclones and climate change. *Nature Geoscience* **3**, 157–163.

11 ———, *et al.*, 2012: Dynamical Downscaling Projections of Late 21st Century Atlantic
12 Hurricane Activity: CMIP3 and CMIP5 Model-based Scenarios. *J. Climate*
13 (submitted)

14 Kumar, A., M. Chen, L. Zhang, W. Wang, Y. Xue, C. Wen, L. Marx, B. Huang, 2012:
15 An Analysis of the Nonstationarity in the Bias of Sea Surface Temperature
16 Forecasts for the NCEP Climate Forecast System (CFS) Version 2. *Mon. Wea.*
17 *Rev.*, **140**, 3003–3016. doi: <http://dx.doi.org/10.1175/MWR-D-11-00335.1>

18 Landsea, C. W., and J. A. Knaff, 2000: How much skill was there in forecasting the very
19 strong 1997–98 El Niño? *Bull. Amer. Meteor. Soc.*, **81**, 2107–2119.

20 ———, G.A. Vecchi, L. Bengtsson, and T.R. Knutson, 2009: Impact of Duration
21 Thresholds on Atlantic Tropical Cyclone Counts. *J. Climate*, **23**, 2508-2519

22 LaRow, T. E., Y. K. Lim, D. W. Shin, E. P. Chassignet, and S. Cocke, 2008: Atlantic
23 basin seasonal hurricane simulations. *J. Climate*, **21**, 3191–3206.

1 ———, L. Stefanova, D. W. Shin, and S. Cocke, Seasonal Atlantic tropical cyclone
2 hindcasting/forecasting using two sea surface temperature datasets, *Geophysical*
3 *Research Letters*, **37**, 1-5, doi:10.1029/2009GL041459, 2010.

4 Latif, M., N. Keenlyside, and J. Bader, 2007: Tropical sea surface temperature, vertical
5 wind shear, and hurricane development. *Geophysical Research Letters*, **34**,
6 L01710, doi:10.1029/2006GL027969.

7 Li, S., and R. Lund, Multiple changepoint detection via genetic algorithms, 2012: *J.*
8 *Climate*, **25**, 674-686.

9 MacAdie, C.J., C.W. Landsea, C.J. Neumann, J.E. David, E. Blake, and G.R. Hammer,
10 2009: *Tropical cyclones of the North Atlantic Ocean, 1851-2006*, Technical
11 Memo, National Climatic Data Center in cooperation with the TCP/National
12 Hurricane Center.

13 Mann, M.E., and K.A. Emanuel, 2006: Atlantic hurricane trends linked to climate
14 change. *Eos, Transactions of the American Geophysical Union*, **87**,
15 doi:10.1029/2006EO240001.

16 Meehl, G., and co-authors, 2012: Decadal Climate Prediction: An Update from the
17 Trenches. *Bull. Amer. Meteorol. Soc.* (submitted).

18 Mendelsohn, R., K. Emanuel, S. Chonabayashi, and L. Bakkensen, 2012: The impact of
19 climate change on global tropical cyclone damage, *Nature Climate Change*, **2**, 205-
20 209.

21 Msadek R., A. Rosati, T. L. Delworth, W. Anderson, G. Vecchi, Y.-S. Chang, K. Dixon,
22 R. G. Gudgel, W. Stern, A. Wittenberg, X. Yang, F. Zeng, R. Zhang, S. Zhang 2012:

1 Predicting North Atlantic decadal variability in the GFDL coupled system: the 1995
2 climate shift event, in preparation

3 van Oldenborgh, G. J. and Doblas-Reyes, F. J. and Wouters, B. and Hazeleger, W., 2012:
4 Decadal prediction skill in a multi-model ensemble. *Climate Dynamics*, **38**, 1263-
5 1280.

6 Oouchi, K., J. Yoshimura, H. Yoshimura, R. Mizuta, S. Kusumoki, and A. Noda, 2006:
7 Tropical cyclone climatology in a global warming climate as simulated in a 20-km-
8 mesh global atmospheric model: Frequency and wind intensity analysis. *Journal of*
9 *the Meteorological Society of Japan* **84**, 259–276.

10 Peduzzi, P., B. Chatenoux, H. Dao, A. De Bono, C. Herold, J. Kossin, F. Mouton, and O.
11 Nordbeck, Global trends in tropical cyclone risk, *Nature Climate Change*, **2**, 289-
12 294, 2012.

13 Pielke, R. A. Jr and coauthors, Normalized hurricane damages in the United States:
14 1900–2005 *Nat. Hazard. Rev.*, **9**, 29–42, 2008.

15 Pohlmann, H., M. Botzet, M. Latif, A. Roesch, M. Wild, and P. Tschuck, 2004:
16 Estimating the decadal predictability of a coupled AOGCM, *J. Climate*, **17**(22),
17 4463–4472.

18 —, J.H. Jungclaus, A. Köhl, D. Stammer, J. Marotzke, 2009: Initializing Decadal
19 Climate Predictions with the GECCO Oceanic Synthesis: Effects on the North
20 Atlantic. *J. Climate*, **22**, 3926–3938.doi: [10.1175/2009JCLI2535.1](https://doi.org/10.1175/2009JCLI2535.1)

21 Ramsay, H. A., and A. H. Sobel, 2011: Effects of relative and absolute sea surface
22 temperature on tropical cyclone potential intensity using a single-column model.
23 *J. Climate*, **24**, 183–193.

- 1 Rayner, N.A., D.E. Parker, E.B. Horton, C.K. Folland, L.V. Alexander, D.P. Rowell,
2 E.C. Kent, and A. Kaplan, 2003: Global analyses of sea surface temperature, sea
3 ice, and night marine air temperature since the late nineteenth century. *J.*
4 *Geophys. Res.*, **108**, 4407, doi:10.1029/2002JD002670.
- 5 Robson, J., 2011: Understanding the performance of a decadal prediction system. U.
6 Reading Ph.D. Thesis, available at:
7 http://www.met.reading.ac.uk/~swr06jir/thesis/JIR_thesis.pdf
- 8 ———, R. Sutton, K. Lohmann, D. Smith, and M. Palmer, 2012: Causes of the rapid
9 warming of the North Atlantic Ocean in the mid 1990s. *Journal of Climate*, **25**,
10 4116-4134.
- 11 Rosati, A. and co-authors, 2012: Decadal Climate Prediction Experiments at GFDL. *J.*
12 *Climate* (submitted).
- 13 Rotstayn, L. D., U Lohmann, 2002: Tropical Rainfall Trends and the Indirect Aerosol
14 Effect. *J. Climate*, **15**, 2103–2116. doi: 10.1175/1520-0442.
- 15 Shen, W., R. E. Tuleya, and I. Ginis, 2000: A sensitivity study of the thermodynamic
16 environment on GFDL model hurricane intensity: Implications for global
17 warming, *Journal of Climate*, **13**, 109-121.
- 18 Smith, D. M., Smith, D., and J. Murphy, 2007: An objective ocean temperature and
19 salinity analysis using covariances from a global climate model, *Journal of*
20 *Geophysical Research*, **112**, doi:10.1029/2005JC003172.
- 21 ———, S. Cusack, A. W. Colman, C. K. Folland, G. R. Harris, and J. M. Murphy 2007:
22 Improved Surface Temperature Prediction for the Coming Decade from a Global
23 Climate Model, *Science*, **317**, 796–799.

1 ———, R. Eade, N.J. Dunstone, D. Fereday, J. M. Murphy, H. Pohlmann, and A.A. Scaife,
2 2010: Skillful multi-year predictions of Atlantic hurricane frequency, *Nature*
3 *Geoscience*, **3**, 846-849.

4 Smith, T.M., R.W. Reynolds, T.C. Peterson, and J. Lawrimore, 2008: Improvement to
5 NOAA's historical merged land-ocean surface temperature analysis (1880–2006).
6 *J. Climate*, **21**, 2283–2296.

7 Sobel, A. H., and C. S. Bretherton, 2000: Modeling tropical precipitation in a single
8 column. *J. Climate*, **13**, 4378–4392.

9 Sobel, A.H., I.M. Held, and C.S. Bretherton, 2002: The ENSO signal in tropical
10 tropospheric temperature. *J. Climate*, **15**, 2702–2706.

11 Stockdale, T.N., 1997: Coupled ocean-atmosphere forecasts in the presence of climate
12 drift, *Mon. Wea. Rev.*, **125**, 809–818.

13 Sugi, M., H. Murakami, and J. Yoshimura, 2009: A reduction in global tropical cyclone
14 frequency due to global warming. *SOLA*, **5**, 164–167.

15 ———, ———, and ———, 2012: On the mechanism of tropical cyclone frequency changes
16 due to global warming. *J. Meteorol. Soc. Japan*, **90A**, 397-408.

17 Sutton, R.T. and D.L.R. Hodson, 2005: Atlantic Ocean forcing of North American and
18 European summer climate, *Science*, **309**(5731), 115-118.

19 Swanson, K.L., 2008: Nonlocality of Atlantic tropical cyclone intensities. *Geochemistry*
20 *Geophysics Geosystems* **9**, Q04V01, doi:10.1029/ 2007GC00184.

21 Tang, B.H., and J.D. Neelin, 2004: ENSO influence on Atlantic hurricanes via
22 tropospheric warming, *Geophysical Research Letters*, **31** (L24204),
23 doi:10.1029/2004GL021072.

1 Taylor, K.E., R.J. Stouffer, and G.A. Meehl, 2012: An overview of CMIP5 and the
2 experiment design. *Bulletin of the American Meteorological Society*, **93**, 485-498.

3 Teng, H., G. Branstator, and G. A. Meehl, 2011: Predictability of the Atlantic overturning
4 circulation and associated surface patterns in two CCSM3 climate change
5 ensemble experiments. *J. Climate*, **24**, 6054-6076.

6 Vecchi, G.A., A.T. Wittenberg and A. Rosati, 2006: Reassessing the role of stochastic
7 forcing in the 1997-8 El Niño. *Geophys. Res. Lett.*, **33**, L01706,
8 doi:10.1029/2005GL024738.

9 ———, and B.J. Soden, 2007a: Effect of remote sea surface temperature change on tropical
10 cyclone potential intensity. *Nature*, **450**, 1066–1071.

11 ———, and ———, 2007b: Global warming and the weakening of the tropical circulation. *J.*
12 *Climate*, **20**(17), 4316-4340.

13 ———, and T.R. Knutson, 2008: On estimates of historical North Atlantic tropical cyclone
14 activity. *J. Climate*, **21**(14), 3580-3600.

15 ———, and ———, 2011: Estimating annual numbers of Atlantic hurricanes missing from
16 the HURDAT database (1878-1965) using ship track density. *J. Climate*, **24**(6),
17 1736-1746

18 ———, K.L. Swanson, and B.J. Soden, 2008: Whither Hurricane Activity? *Science* **322**
19 (5902), 687-689.

20 ———, M. Zhao, H. Wang, G. Villarini, A. Rosati, A. Kumar, I. M. Held, and R. Gudgel,
21 2011: Statistical-dynamical predictions of seasonal North Atlantic hurricane
22 activity, *Monthly Weather Review*, **139**(4), 1070-1082.

- 1 —, S. Fueglistaler, I.M. Held, T.R. Knutson, and M. Zhao, 2012: Impacts of
2 Atmospheric Temperature Changes on Tropical Cyclone Activity. *J. Climate*
3 (submitted).
- 4 Villarini, G., and G.A. Vecchi. 2012a: North Atlantic Power Dissipation Index (PDI) and
5 Accumulated Cyclone Energy (ACE): Statistical modeling and sensitivity to sea
6 surface temperature changes. *Journal of Climate* **25**(2), 625-637.
- 7 —, and —, 2012b: Twenty-first-century projections of North Atlantic tropical
8 storms from CMIP5 models, *Nature Climate Change*,
9 doi:10.1038/NCLIMATE1530.
- 10 —, and —, 2012c: Projected increases in North Atlantic tropical cyclone intensity
11 from CMIP5 models, *J. Climate* doi:10.1175/JCLI-D-12-00441.
- 12 —, and —, 2012d: Multi-season lead forecast of the North Atlantic Power
13 Dissipation Index (PDI) and Accumulated Cyclone Energy (ACE). *J. Climate*
14 doi:10.1175/JCLI-D-12-00448..
- 15 —, —, and J.A. Smith, 2010: Modeling of the dependence of tropical storm counts
16 in the North Atlantic Basin on climate indices. *Monthly Weather Review* **138**(7),
17 2681–2705.
- 18 —, —, and —, 2012: U.S. landfalling and North Atlantic hurricanes: Statistical
19 modeling of their frequencies and ratios. *Monthly Weather Review*, 140, 44–65.
- 20 —, —, T.R. Knutson, M. Zhao and J.A. Smith, 2011a: Reconciling differing model
21 projections of changes in the frequency of tropical storms in the North Atlantic basin
22 in a warmer climate, *J. Climate*, **24**(13), 3224-3238.

1 ———, ———, ———, and J.A. Smith, 2011b: Is the recorded increase in short duration North
2 Atlantic tropical storms spurious? *J. Geophys. Res.* **116**, D10114,
3 doi:10.1029/2010JD015493.

4 Vitart, F., Seasonal forecasting of tropical storm frequency using a multi-model
5 ensemble, *Quarterly Journal of the Royal Meteorological Society*, **132**, 647-666,
6 2006.

7 ———, M. Huddleston, D. Deque, T. Palmer, T. Stockdale, M. Davey, S. Ineson, and
8 A. Weisheimer, 2007. Dynamically-based seasonal forecasts of Atlantic tropical
9 storm activity issued in June by EUROSIP, *Geophysical Research Letters*, **34**
10 (L16815), doi:10.1029/2007GL030740.

11 Von Storch, H., and F.W. Zwiers, 1999: *Statistical Analysis in Climate Research*,
12 Cambridge University Press, 484 pp.

13 Wang, H., J.K.E. Schemm, A. Kumar, W. Wang, L. Long, M. Chelliah, G.D. Bell, and P.
14 Peng, 2009: A statistical forecast model for Atlantic seasonal hurricane activity
15 based on the NCEP dynamical seasonal forecast, *Journal of Climate*, **22**, 4481-
16 4500.

17 Yang, X. and co-authors (2012): A predictable AMO-like pattern in GFDL's fully-
18 coupled ensemble initialization and decadal forecasting system. *J. Climate*,
19 doi:10.1175/JCLI-D-12-00231

20 Yeager, S., A. Karspeck, G. Danabasoglu, J. Tribbia, and H. Teng, 2012: A Decadal
21 Prediction Case Study: Late Twentieth-Century North Atlantic Ocean Heat Content.
22 *J. Climate*, **25**, 5173–5189. doi:10.1175/JCLI-D-11-00595.1

1 Zhang, R., and T.L. Delworth, 2005: Simulated tropical response to a substantial
2 weakening of the Atlantic thermohaline circulation. *Journal of Climate*, **18**, 1853-
3 1860.

4 ———, and ———, 2006: Impact of Atlantic multidecadal oscillations on India/Sahel rainfall
5 and Atlantic hurricanes. *Geophysical Research Letters*, **33**, L17712,
6 doi:10.1029/2006GL026267.

7 ———, and ———, 2009: A new method for attributing climate variations over the Atlantic
8 hurricane basin's main development region. *Geophysical Research Letters*, **36**,
9 L06701, doi:10.1029/2009GL037260.

10 ———, and coauthors, 2012: Have aerosols caused the observed Atlantic Multidecadal
11 Variability? *Nature*, submitted.

12 Zhang, S., and A. Rosati, 2010: An inflated ensemble filter for ocean data assimilation
13 with a biased coupled GCM. *Mon. Wea. Rev.*, **138**(10), 3905-3931.

14 ———, M.J. Harrison, A. Rosati, and A.T. Wittenberg, 2007: System design and
15 evaluation of coupled ensemble data assimilation for global oceanic climate
16 studies. *Mon. Wea. Rev.*, **135**, 3541–3564.

17 Zhao, M., and I.M. Held, 2011: The response of tropical cyclone statistics to an increase
18 in CO₂ with fixed sea surface temperatures. *J. Climate*, **24**, 5353–5364.

19 ———, ———, S.-J. Lin, and G.A. Vecchi, 2009: Simulations of global hurricane
20 climatology, interannual variability, and response to global warming using a 50-
21 km resolution GCM. *J. Climate*, **22**, 6653–6678.

1 ———, ———, and G.A. Vecchi, 2010: Retrospective forecasts of the hurricane season using
2 a global atmospheric model assuming persistence of SST anomalies. *Mon. Wea.*
3 *Rev.*, **138**, 3858–3868.
4

Figure Captions:

Figure 1: Retrospective and future forecasts of hurricane frequency. Upper panels show the retrospective forecasts for five-year running hurricane frequency, lower panels focus on the nine-year running forecasts. Left panels show the results from uninitialized experiments, while the right panels show the results for initialized experiments. Black line shows the observed five-year (upper) and nine-year (lower) hurricane counts from the NOAA Hurricane Database (HURDAT; Jarvinen *et al.* 1984, MacAdie *et al.* 2009) that includes an adjustment for observing inhomogeneity prior to 1966 described in Vecchi and Knutson (2011). Retrospective forecasts are shown in: red line shows the forecasts from the GFDL-CM2.1 system, blue line shows the UKMO-DePreSys-PPE system, and the yellow line shows the two-system ensemble-mean.

Figure 2: Correlation for retrospective multi-year forecasts of North Atlantic hurricane frequency, with 90% uncertainty estimates. Each cluster of bars shows the retrospective correlation of multi-year hurricane frequency forecasts for Lead 2-6 years (left), Lead 6-10 years (middle) and Lead 2-10 years (right). Gray symbol is the correlation of the persistence of the five-year average count preceding the initialization of the model. Red symbols are for the GFDL-DecPre system, blue are for UKMO-DePreSys-PPE, and yellow is for the two system average. The initialized and uninitialized versions of each model are distinguished by different coloring. The sample correlation estimate is shown by the circle, the bars show the two-sided 90% uncertainty of a correlation given an underlying correlation with the value shown by the corresponding circle. Asterisk on top of a bar shows correlations that are significantly different from a null hypothesis of an

underlying correlation of zero at $p=0.1$, single-sided, with the effective degrees of freedom estimated as in Bretherton *et al.* (1999).

Figure 3: Mean Skill Score Squared (MSSS) of retrospective initialized multi-year hurricane frequency forecasts for various leads and models. Horizontal axis shows the MSSS against climatology, vertical axis shows the MSSS against the uninitialized forecasts; diagonal line indicates the one-to-one line. Left panel shows MSSS values for the five-year running-mean forecasts, right panel shows MSSS values for the nine-year running-mean forecasts. Circles show the values for the GFDL-DecPre system, squares for UKMO-DePreSys-PPE, and stars for the two-model ensemble mean. Different colors indicate different forecast leads.

Figure 4: Retrospective and future forecasts of the SST indices used for the hurricane emulator. Left panels show time-series of the five-year mean SSTA anomalies averaged over the global tropics (upper) and Atlantic hurricane main development region (lower), at lead 2-6. Black lines show observational estimates from HadISST.v1 (Rayner *et al.* 2003; solid) and ERSST.v3b (Smith *et al.* 2008; dotted). Colored lines show initialized (dashed) and uninitialized (solid) experiments from GFDL-DecPre (reds) and UKMO-DePreSys-PPE (blue). Right panels show the retrospective correlations of the forecasts at Lead 2-6 against the HadISST.v1 SST product.

Figure 5: Retrospective forecasts exploring the source of the initialized vs. uninitialized components. Left panel takes Atlantic MDR SST from initialized experiments and

tropical-mean SST from uninitialized, right panel takes tropical-mean SST from initialized experiments and Atlantic MDR SST from uninitialized experiments. The skill comes from the improvement of tropical Atlantic SST in the initialized experiments.

Figure 6: Retrospective forecasts arranged by verification and initialization date. Top panels (a and b) show the retrospective forecasts of five-year running hurricane averages for various leads, arranged so that each point on the time axis corresponds to the midpoint of the five-year interval over which the average is computed (*e.g.*, 1992 corresponds to the midpoint of the 1990-1994 average). Bottom panels (c and d) show the retrospective five-year forecasts for various leads arranged so that each point on the time axis corresponds to the date in which the model was initialized. Left panels are from the GFDL-CM2.1 forecasts, right panels are from the UKMO-DePreSys-PPE system. Dark line in the top panels shows the observed five-year running counts.

Figure 7: Empirical probability density function (PDF) estimates for the change in seasonal hurricane counts over the entire record and over the four years that preceded the 1994-1995 climate shift. The quantity explored is the difference in hurricane counts averaged over the five years following a given year with the counts of that year (*e.g.*, for 1991 it is the difference of hurricane counts averaged 1992-1996 with those in 1991); PDFs are estimated through Gaussian convolution with an *e*-folding scale of 2.5 hurricanes per year. Black lines are based on observations, blue lines on the forecasts with GFDL-DecPre, and red lines on the forecasts using UKMO-DePreSys; solid lines are computed over the 1961-2006 period, dashed lines over 1991-1994. The separation of

the solid and dashed black line is a reflection of the increase in storm counts that occurred in 1995. Notice that there is no tendency for forecasts initialized in the early-1990s to have indicate a tendency for intensification through the early years of the forecast: the forecast systems do not dynamically predict the occurrence of the 1994-1995 shift.

Figure 8: Retrospective forecasts of North Atlantic hurricane frequency after removing 1994-1995 shift in the mean from forecasts and verification (see Section III.A). Left panel shows the initialized forecasts at lead 2-6, right panels show the uninitialized experiments. Black line shows the observed counts, red line is from the GFDL-DecPre system, blue line is from UKMO-DePreSys-PPE and the yellow line is the two system average, all after removing the 1994-1995 shift in the mean.

Figure 9: Retrospective correlations of forecasts after removing 1994-1995 shift in the mean from forecasts and verification. Gray symbol is the correlation of the persistence of the five-year average count preceding the initialization of the model. Red symbols are for the GFDL-DecPre system, blue are for UKMO-DePreSys-PPE, and yellow is for the two system average. The initialized and uninitialized versions of each model are distinguished by different coloring. The sample correlation estimate is shown by the circle, the bars show the two-sided 90% uncertainty of a correlation given an underlying correlation with the value shown by the corresponding circle. Asterisk on top of a bar shows correlations that are significantly different from a null hypothesis of an underlying correlation of zero at $p=0.1$, single-sided, with the effective degrees of freedom estimated as in Bretherton *et al.* (1999).

1 **Figure 10:** Impact of Argo on retrospective and future forecasts of hurricane frequency
2 using GFDL-DecPre. Lagged-ensemble (Lead 1-5 & Lead 2-6) forecasts of five-year
3 Atlantic hurricane frequency based on the standard GFDL-DecPre system (gray line), and
4 from a perturbation experiment in which forecasts initialized 2004 and later do not
5 include data from Argo floats in the initialization (dashed line); black line shows
6 observed five-year counts. A change in the drift of the initialized forecasts after the
7 introduction of Argo leads to an increase in the predicted number of hurricanes after
8 2004.
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Forecast system	Underlying GCM	Initialization Procedure	Ensemble Type	Initialization Date	Treatment of Volcanoes
GFDL-CM2.1 DecPre (Rosati <i>et al.</i> 2012; Yang <i>et al.</i> 2012)	GFDL-CM2.1 (Delworth <i>et al.</i> 2006)	Fully Coupled Ensemble Kalman Filter (Zhang <i>et al.</i> 2007), full variable assimilation	Ten ensemble members from the EnKF assimilation	1-January of each year 1960-2011.	Future volcanoes included in radiative forcing
UKMO DepPreSys-PPE (Smith <i>et al.</i> 2007; Smith <i>et al.</i> 2010)	HadCM3 (Gordon <i>et al.</i> 2000)	Atmospheric and oceanic conditions relaxed to observations. Ocean anomaly initialization. (Smith and Murphy 2007)	Nine ensemble member perturbed physics ensemble (PPE)	1 November of each year 1960-2005.	Forcing from past volcanic forcing included

4 **Table 1:** Summary of the two dynamical multi-year experimental forecast systems
5 explored in this manuscript.

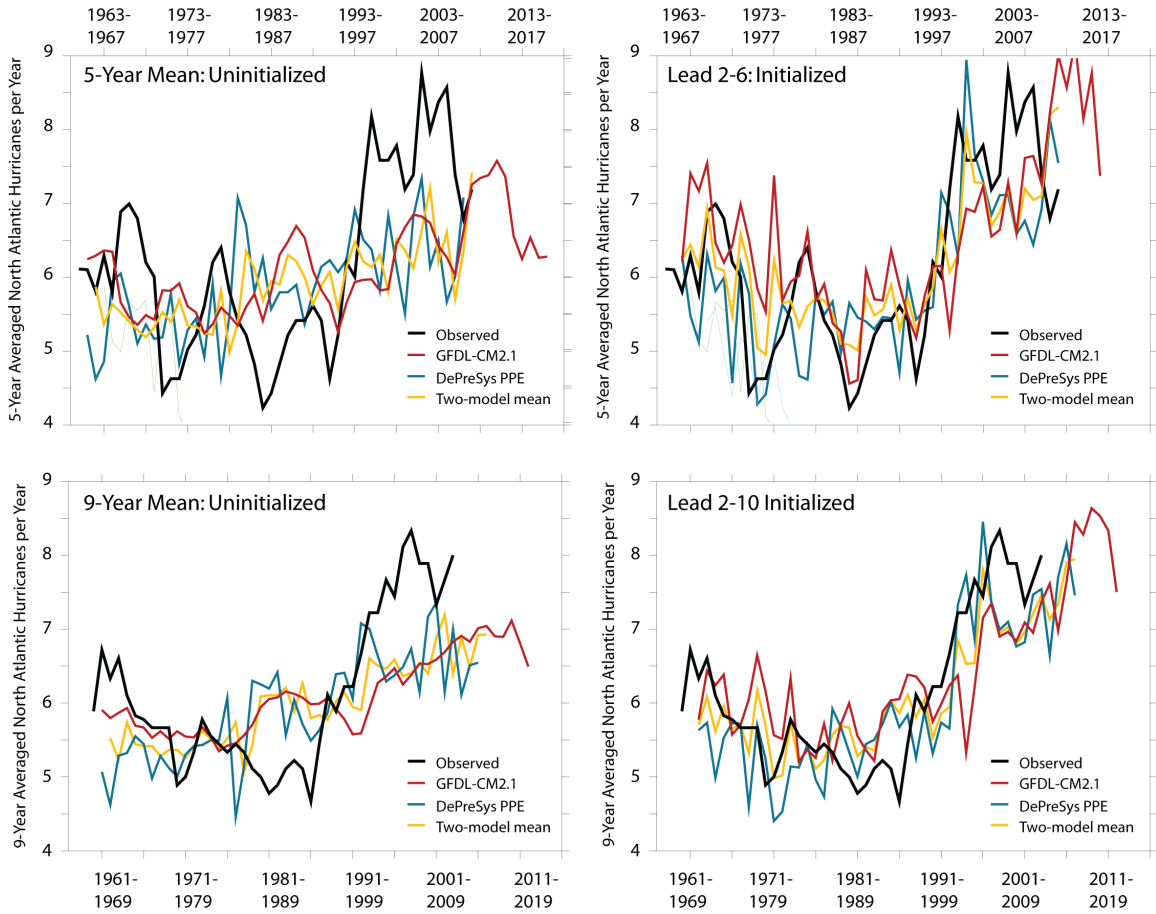


Figure 1: Retrospective and future forecasts of hurricane frequency. Upper panels show the retrospective forecasts for five-year running hurricane frequency, lower panels focus on the nine-year running forecasts. Left panels show the results from uninitialized experiments, while the right panels show the results for initialized experiments. Black line shows the observed five-year hurricane counts from the NOAA Hurricane Database (HURDAT; Jarvinen *et al.* 1984, MacAdie *et al.* 2009) that includes an adjustment for observing inhomogeneity prior to 1966 described in Vecchi and Knutson (2011). Retrospective forecasts are shown in: red line shows the forecasts from the GFDL-CM2.1 system, blue line shows the UKMO-DePreSys-PPE system, and the yellow line shows the two-system ensemble-mean.

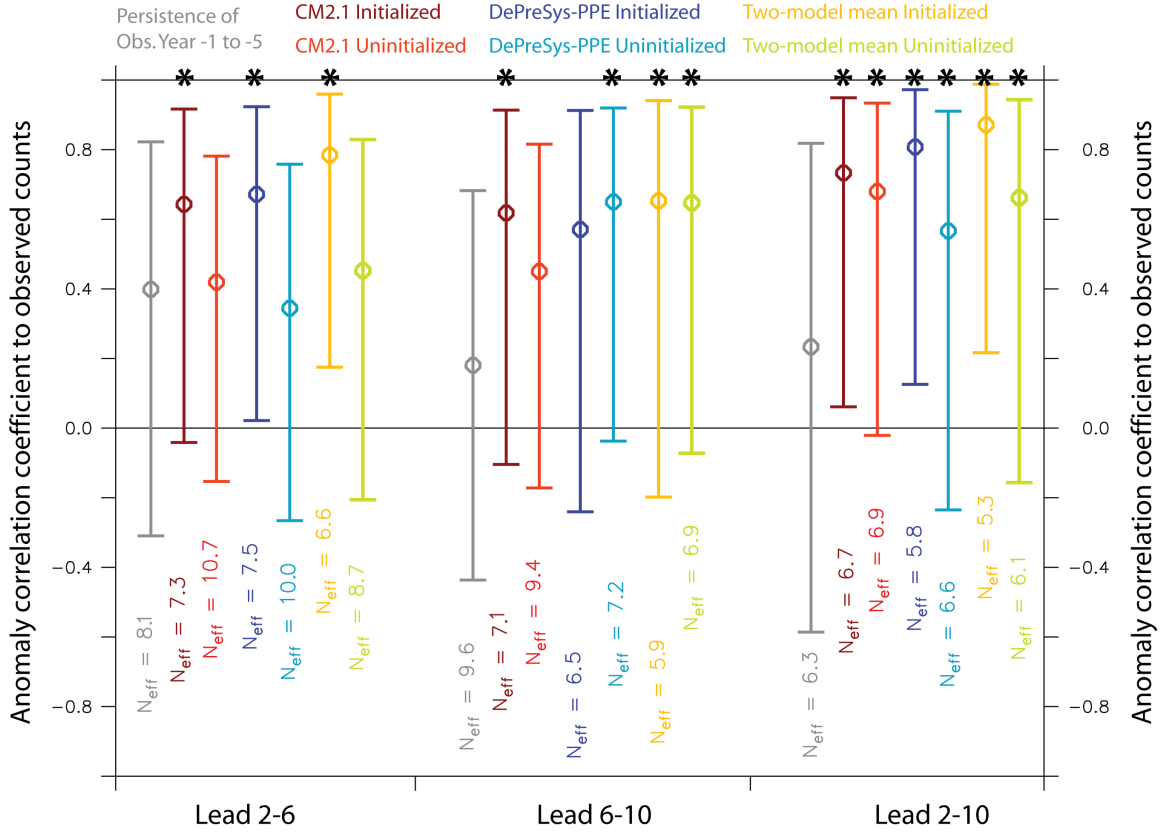


Figure 2: Correlation for retrospective multi-year forecasts of North Atlantic hurricane frequency, with 90% uncertainty estimates. Each cluster of bars shows the retrospective correlation of multi-year hurricane frequency forecasts for lead 2-6 years (left), lead 6-10 years (middle) and lead 2-10 years (right). Gray symbol is the correlation of the persistence of the five-year average count preceding the initialization of the model. Red symbols are for the GFDL-DecPre system, blue are for UKMO-DePreSys-PPE, and yellow is for the two system average. The initialized and uninitialized versions of each model are distinguished by different coloring. The sample correlation estimate is shown by the circle, the bars show the two-sided 90% uncertainty of a correlation given an underlying correlation with the value shown by the corresponding circle. Asterisk on top of a bar shows correlations that are significantly different from a null hypothesis of an underlying correlation of zero at $p=0.1$, single-sided, with the effective degrees of freedom estimated as in Bretherton *et al.* (1999).

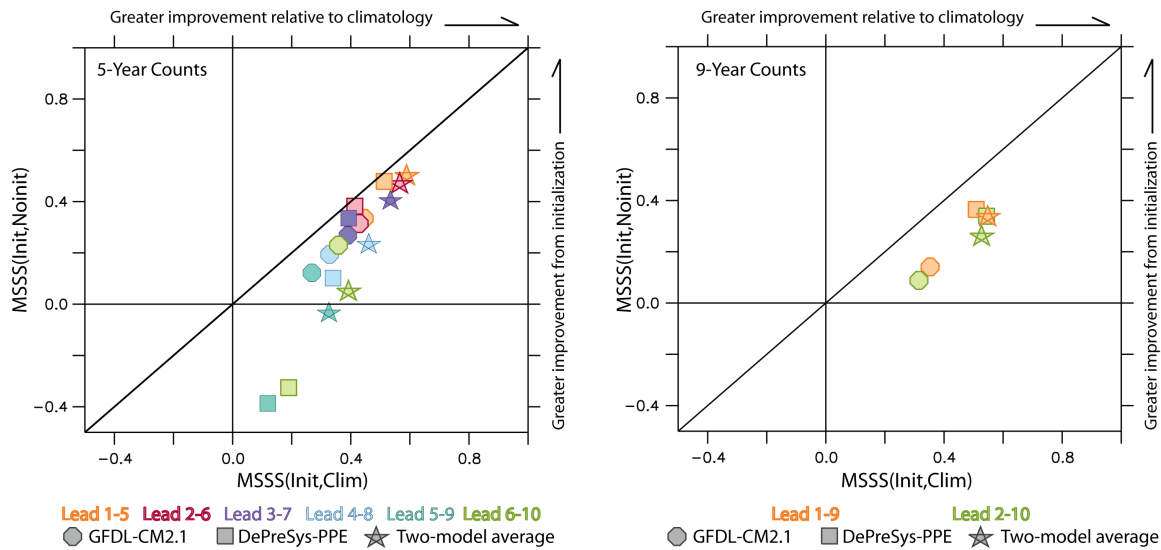


Figure 3: Mean Skill Score Squared (MSSS) of retrospective initialized multi-year hurricane frequency forecasts for various leads and models. Horizontal axis shows the MSSS against climatology, vertical axis shows the MSSS against the uninitialized forecasts; diagonal line indicates the one-to-one line. Left panel shows MSSS values for the five-year running-mean forecasts, right panel shows MSSS values for the nine-year running-mean forecasts. Circles show the values for the GFDL-DecPre system, squares for UKMO-DePreSys-PPE, and stars for the two-model ensemble mean. Different colors indicate different forecast leads.

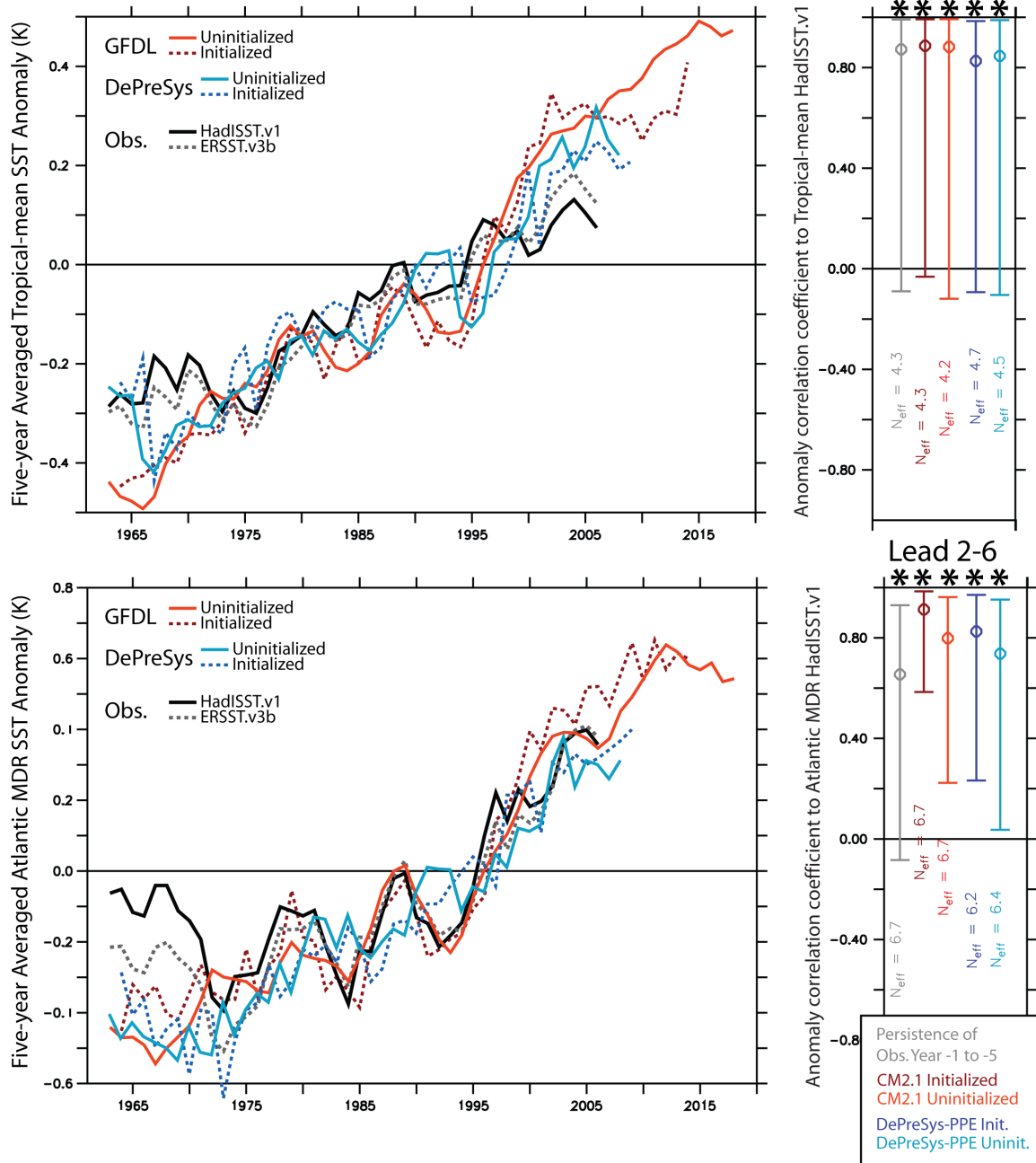
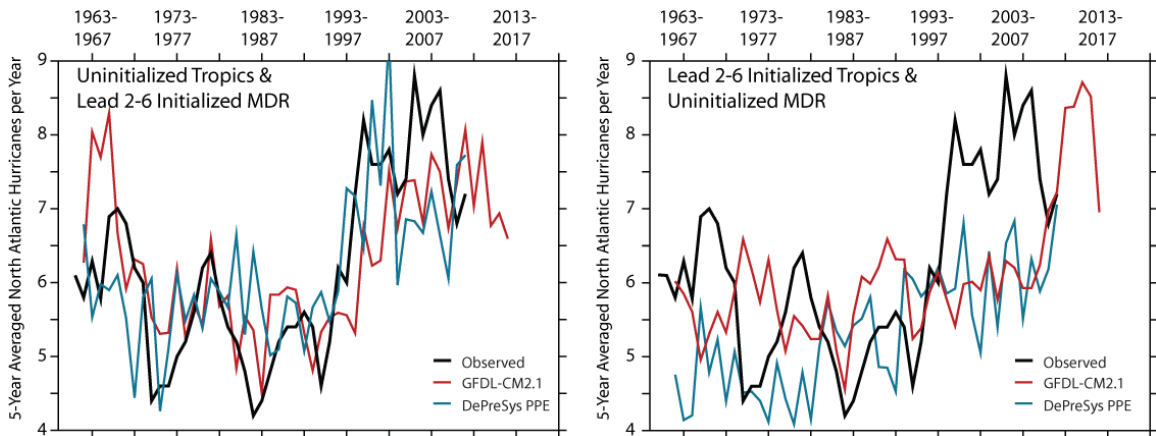


Figure 4: Retrospective and future forecasts of the SST indices used for the hurricane emulator. Left panels show time-series of the five-year mean SST anomalies averaged over the global tropics (upper) and Atlantic hurricane main development region (lower), at lead 2-6. Black lines show observational estimates from HadISST.v1 (Rayner *et al.* 2003; solid) and ERSST.v3b (Smith *et al.* 2008; dotted). Colored lines show initialized (dashed) and uninitialized (solid) experiments from GFDL-DecPre (reds) and UKMO-DePreSys-PPE (blue). Right panels show the retrospective correlations of the forecasts at lead 2-6 against the HadISST.v1 SST product.

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Figure 5: Retrospective forecasts exploring the source of the initialized vs. uninitialized components. Left panel takes Atlantic MDR SST from initialized experiments and tropical-mean SST from uninitialized experiments, right panel takes tropical-mean SST from initialized experiments and Atlantic MDR SST from uninitialized experiments. The skill comes from the improvement of tropical Atlantic SST in the initialized experiments.

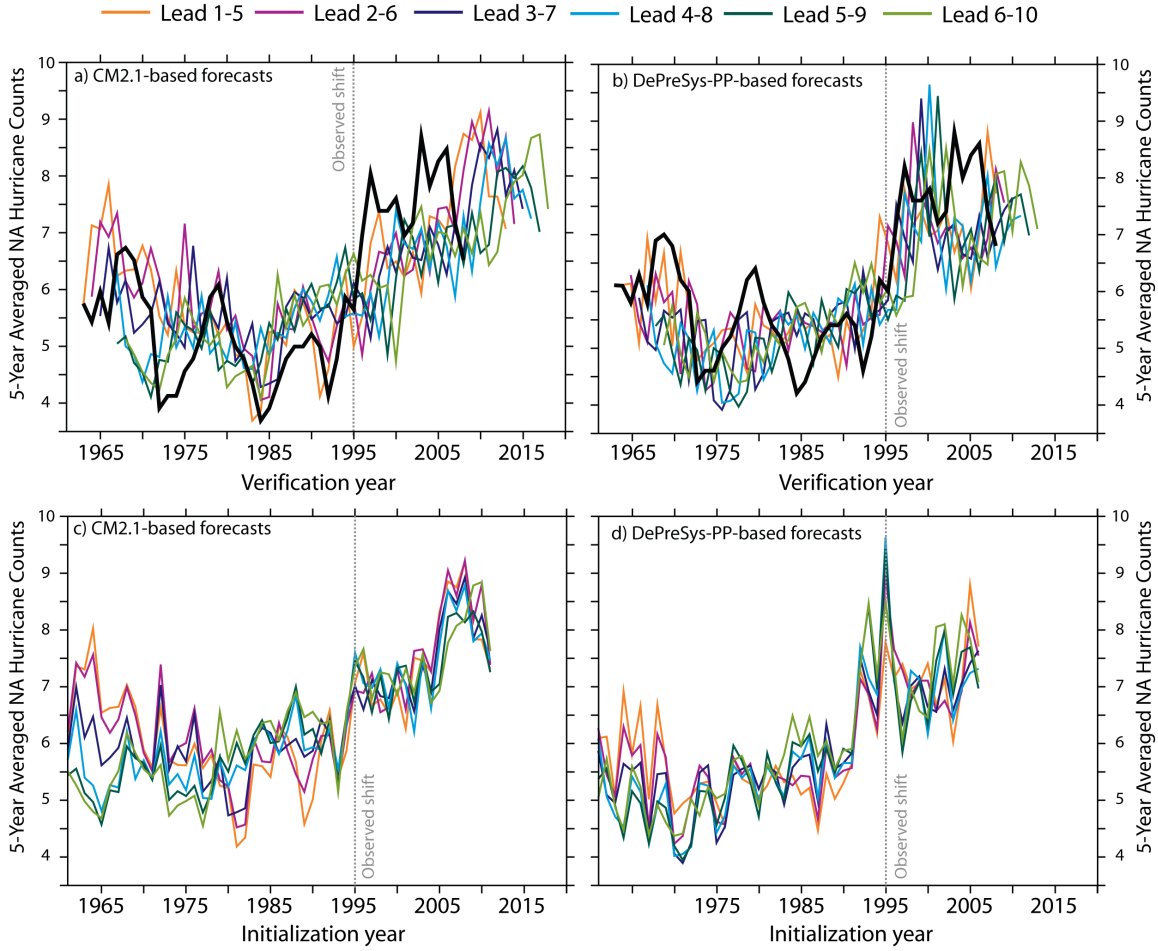
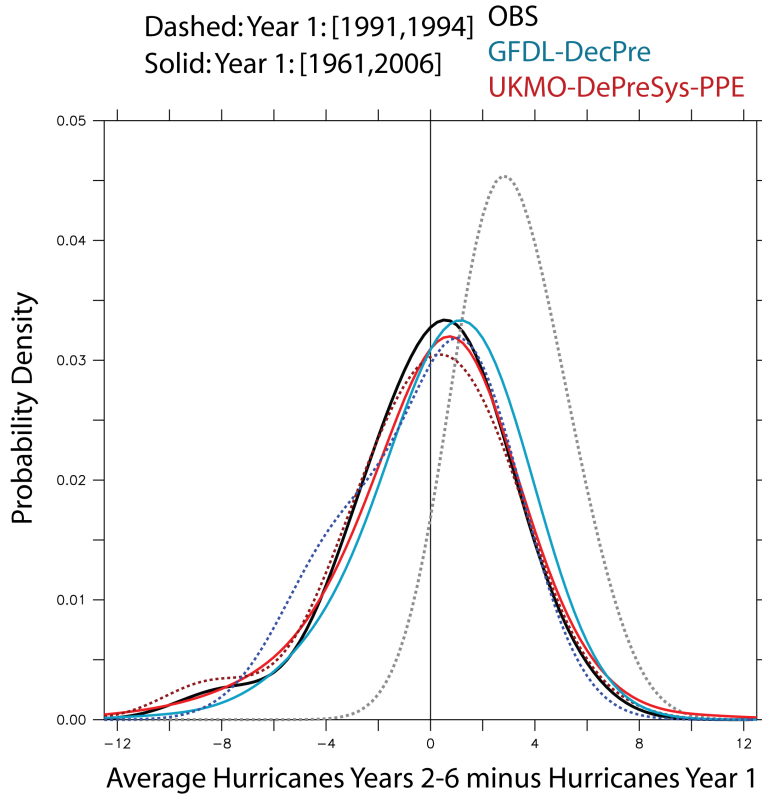


Figure 6: Retrospective forecasts arranged by verification and initialization date. Top panels (a and b) show the retrospective forecasts of five-year running hurricane averages for various leads, arranged so that each point on the time axis corresponds to the midpoint of the five-year interval over which the average is computed (*e.g.*, 1992 corresponds to the midpoint of the 1990-1994 average). Bottom panels (c and d) show the retrospective five-year forecasts for various leads arranged so that each point on the time axis corresponds to the date in which the model was initialized. Left panels are from the GFDL-CM2.1 forecasts, right panels are from the UKMO-DePreSys-PPE system. Dark line in the top panels shows the observed five-year running counts.



1 Average Hurricanes Years 2-6 minus Hurricanes Year 1
 2 **Figure 7:** Empirical probability density function (PDF) estimates for the change in
 3 seasonal hurricane counts over the entire record and over the four years that preceded the
 4 1994-1995 climate shift. The quantity explored is the difference in hurricane counts
 5 averaged over the five years following a given year with the counts of that year (*e.g.*, for
 6 1991 it is the difference of hurricane counts averaged 1992-1996 with those in 1991);
 7 PDFs are estimated through Gaussian convolution with an e -folding scale of 2.5
 8 hurricanes per year. Black lines are based on observations, blue lines on the forecasts
 9 with GFDL-DecPre, and red lines on the forecasts using UKMO-DePreSys; solid lines
 10 are computed over the 1961-2006 period, dashed lines over 1991-1994. PDFs of the
 11 models are based on the various ensemble members. The separation of the solid and
 12 dashed black lines is a reflection of the increase in storm counts that occurred in 1995.
 13 Notice that there is no tendency for forecasts initialized in the early-1990s to have
 14 indicate a tendency for frequency increase through the early years of the forecast: the
 15 forecast systems do not dynamically predict the occurrence of the 1994-1995 shift.

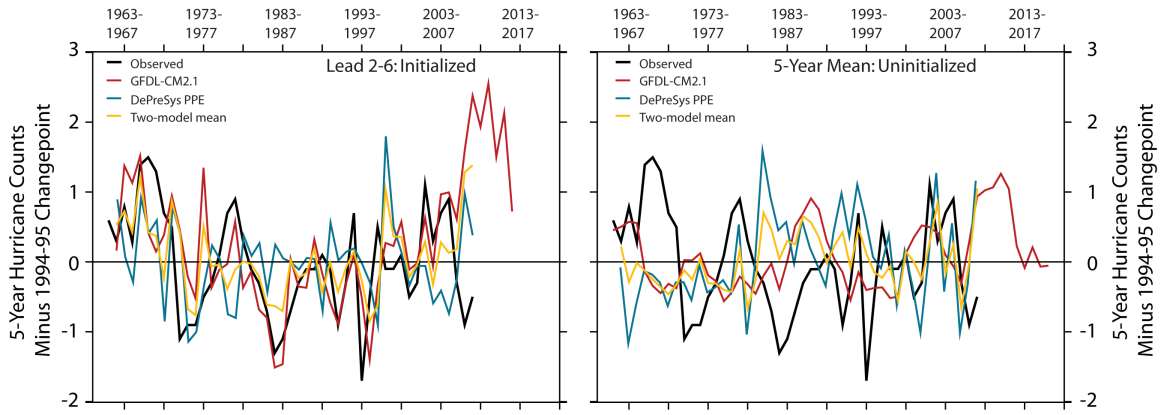


Figure 8: Retrospective forecasts of North Atlantic hurricane frequency after removing 1994-1995 shift in the mean from forecasts and verification (see Section III.A). Left panel shows the initialized forecasts at lead 2-6, right panel shows the uninitialized experiments. Black line shows the observed counts, red line is from the GFDL-DecPre system, blue line is from UKMO-DePreSys-PPE and the yellow line is the two system average, all after removing the 1994-1995 shift in the mean.

Persistence of Obs. Year -1 to -5 CM2.1 Initialized DePreSys-PPE Initialized Two-model mean Initialized
 CM2.1 Uninitialized DePreSys-PPE Uninitialized Two-model mean Uninitialized

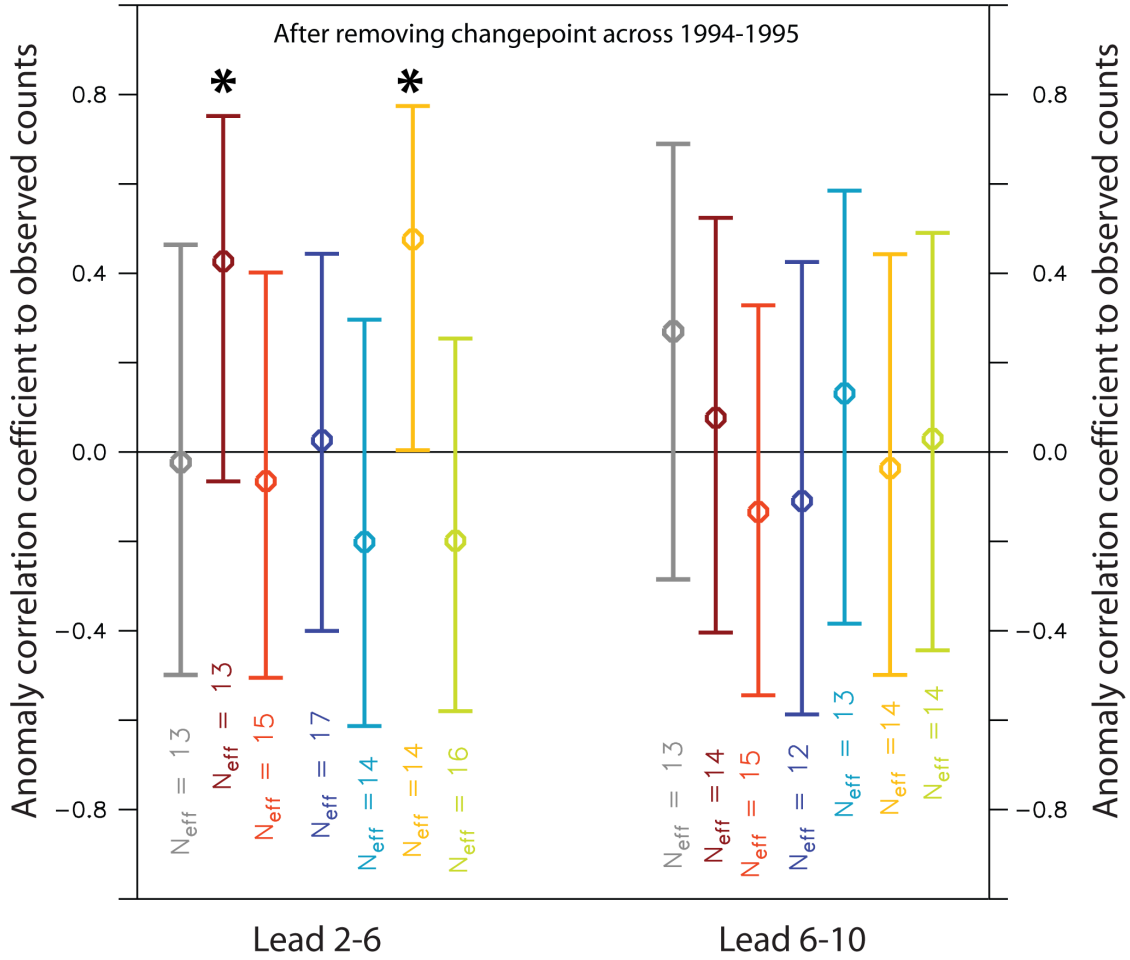


Figure 9: Retrospective correlations of forecasts after removing 1994-1995 shift in the mean from forecasts and verification. Gray symbol is the correlation of the persistence of the five-year average count preceding the initialization of the model. Red symbols are for the GFDL-DecPre system, blue are for UKMO-DePreSys-PPE, and yellow is for the two system average. The initialized and uninitialized versions of each model are distinguished by different coloring. The sample correlation estimate is shown by the circle, the bars show the two-sided 90% uncertainty of a correlation given an underlying correlation with the value shown by the corresponding circle. Asterisk on top of a bar shows correlations that are significantly different from a null hypothesis of an underlying correlation of zero at $p=0.1$, single-sided, with the effective degrees of freedom estimated as in Bretherton *et al.* (1999).

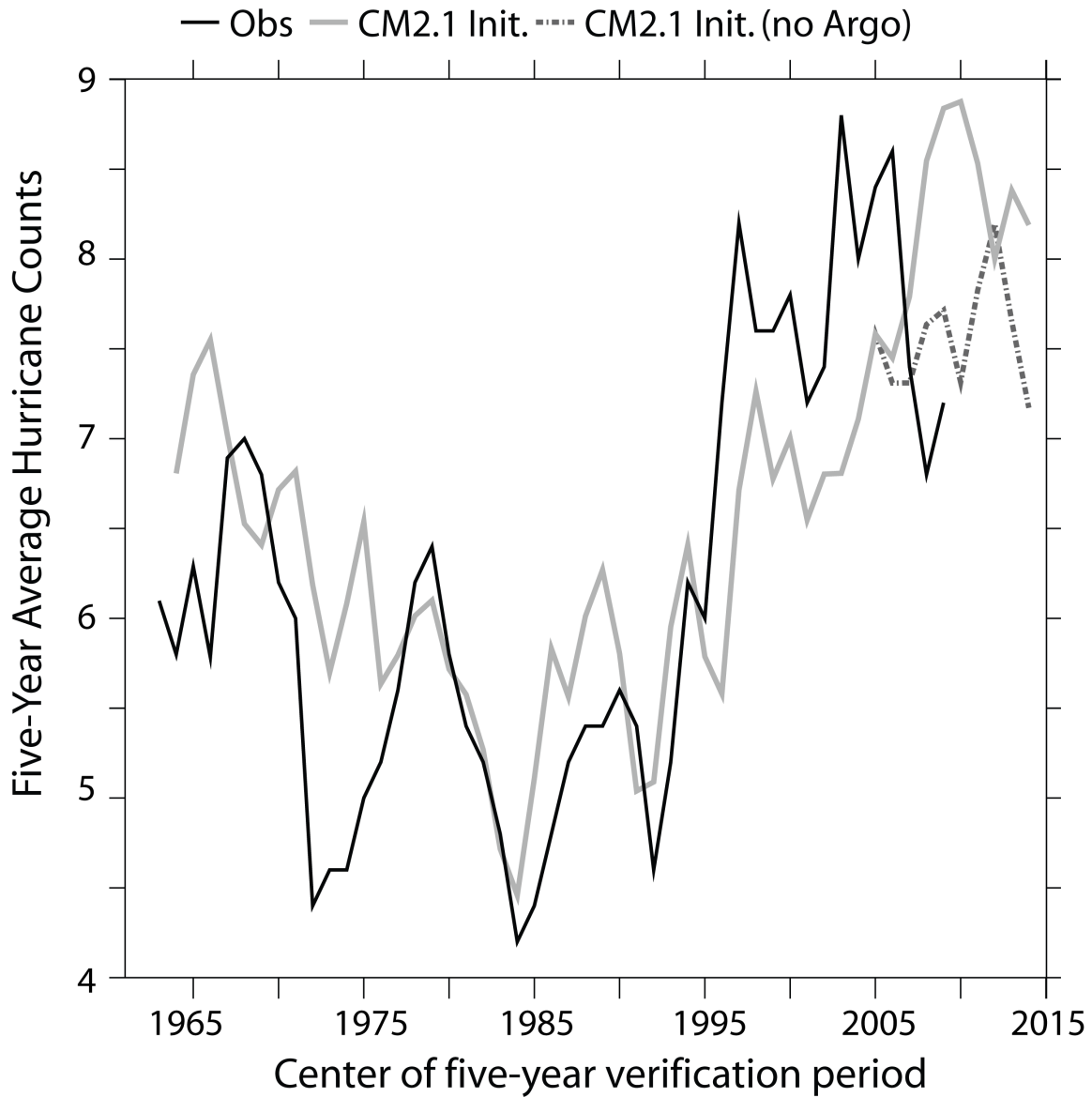


Figure 10: Impact of Argo on retrospective and future forecasts of hurricane frequency using GFDL-DecPre. Lagged-ensemble (Lead 1-5 & Lead 2-6) forecasts of five-year Atlantic hurricane frequency based on the standard GFDL-DecPre system (gray line), and from a perturbation experiment in which forecasts initialized 2004 and later do not include data from Argo floats in the initialization (dashed line); black line shows observed five-year counts. A change in the drift of the initialized forecasts after the introduction of Argo leads to an increase in the predicted number of hurricanes after 2004.